

Current and emerging developments in subseasonal to decadal prediction

Article

Accepted Version

Merryfield, W. J., Baehr, J., Batte, L., Becker, E. J., Butler, A. H., Coelho, C. A., Danabasoglu, G., Dirmeyer, P. A., Doblas-Reyes, F. J., Domeisen, D. I., Ferranti, L., Ilynia, T., Kumar, A., Muller, W. A., Rixen, M., Robertson, A. W., Smith, D. M., Takaya, Y., Tuma, M., Vitart, F., White, C. J., Alvarez, M. S., Ardilouze, C., Attard, H., Baggett, C., Balmaseda, M. A., Beraki, A. F., Battcharjee, P. S., Bilbao, R., Marques De Andrade, F., DeFlorio, M. J., Diaz, L. B., Ehsan, M. A., Frangkoulidis, G., Grainger, S., Green, B. W., Hell, M. C., Infanti, J. M., Isensee, K., Kataoka, T., Kirtman, B. P., Klingaman, N. P., Lee, J.-Y., Mayer, K., McKay, R., Mecking, J., Miller, D. E., Neddermann, N., Ng, C. H., Osso, A., Pankatz, K., Peatman, S., Pegion, K., Perwitz, J., Raclade-Coronel, G. C., Reintges, A., Renkl, C., Solaraju-Murali, B., Spring, A., Stan, C., Sun, Y. Q., Tozer, C. R., Vigaud, N., Woolnough, S. and Yeager, S. (2020) Current and emerging developments in subseasonal to decadal prediction. *Bulletin of the American Meteorological Society*, 101 (6). E869-E896. ISSN 1520-0477 doi: <https://doi.org/10.1175/BAMS-D-19-0037.1> Available at <https://centaur.reading.ac.uk/88780/>

work. See [Guidance on citing](#).

To link to this article DOI: <http://dx.doi.org/10.1175/BAMS-D-19-0037.1>

Publisher: American Meteorological Society

All outputs in CentAUR are protected by Intellectual Property Rights law, including copyright law. Copyright and IPR is retained by the creators or other copyright holders. Terms and conditions for use of this material are defined in the [End User Agreement](#).

www.reading.ac.uk/centaur

CentAUR

Central Archive at the University of Reading

Reading's research outputs online

Current and emerging developments in subseasonal to decadal prediction

William J. Merryfield

*Canadian Centre for Climate Modelling and Analysis, Environment and Climate Change Canada,
Victoria, Canada*

Johanna Baehr

Institute of Oceanography, University of Hamburg, Hamburg, Germany

Lauriane Batté

CNRM, Université de Toulouse, Météo France, CNRS, Toulouse, France

Emily J. Becker

NOAA/NWS/NCEP/Climate Prediction Center/Innovim, LLC, College Park, Maryland

Amy H. Butler

*CIRES, University of Colorado Boulder, and Chemical Sciences Division, NOAA/ESRL, Boulder,
Colorado*

Caio A. S. Coelho

CPTEC/INPE Center for Weather Forecasts and Climate Studies, Cachoeira Paulista, Brazil

Gokhan Danabasoglu

Climate and Global Dynamics Laboratory, NCAR, Boulder, Colorado

Paul A. Dirmeyer

Center for Ocean–Land–Atmosphere Studies, George Mason University, Fairfax, Virginia

Francisco J. Doblas-Reyes

Barcelona Supercomputing Center and ICREA, Barcelona, Spain

Daniela I. V. Domeisen

Institute for Atmospheric and Climate Science, ETH Zürich, Zurich, Switzerland

Laura Ferranti

ECMWF, Reading, United Kingdom

Tatiana Ilynia

Max Planck Institute for Meteorology, Hamburg, Germany

Arun Kumar

Climate Prediction Center, NOAA/NWS/NCEP, College Park, Maryland

44 Wolfgang A. Müller
45 *Max Planck Institute for Meteorology and Deutscher Wetterdienst, Hamburg, Germany*
46
47 Michel Rixen
48 *World Climate Research Programme, World Meteorological Organization, Geneva, Switzerland*
49
50 Andrew W. Robertson
51 *International Research Institute for Climate and Society (IRI), Columbia University, Palisades, NY*
52
53 Doug M. Smith
54 *Met Office Hadley Centre, Met Office, Exeter, UK*
55
56 Yuhei Takaya
57 *Department of Atmosphere, Ocean and Earth System Modeling Research, Meteorological*
58 *Research Institute, Japan Meteorological Agency, Tsukuba, Japan*
59
60 Matthias Tuma
61 *World Climate Research Programme, World Meteorological Organization, Geneva, Switzerland*
62
63 Frederic Vitart
64 *ECMWF, Reading, United Kingdom*
65
66 Christopher J. White
67 *Department of Civil and Environmental Engineering, University of Strathclyde, Glasgow, United*
68 *Kingdom*
69
70 Mariano S. Alvarez
71 *Universidad de Buenos Aires, Centro de Investigaciones del Mar y la Atmósfera, Institut Franco-*
72 *Argentin d'Estudes sur le Climat et ses Impacts, Buenos Aires, Argentina*
73
74 Constantin Ardilouze
75 *CNRM, Université de Toulouse, Météo France, CNRS, Toulouse, France*
76
77 Hannah Attard
78 *Embry-Riddle Aeronautical University-Worldwide, Daytona Beach, Florida*
79
80 Cory Baggett
81 *Department of Atmospheric Science, Colorado State University, Fort Collins, Colorado, and*
82 *NOAA/NWS/NCEP/Climate Prediction Center/Innovim, LLC, College Park, Maryland*
83
84 Magdalena A. Balmaseda
85 *ECMWF, Reading, United Kingdom*
86
87 Asmerom F. Beraki

88 *CSIR – Global Change, Climate and Air Quality Modelling, and Department of Geography,*
89 *Geoinformatics and Meteorology, University of Pretoria, Pretoria, South Africa*
90
91 Partha S. Bhattacharjee
92 *I.M. Systems Group, NOAA/NWS National Centers for Environmental Prediction, College Park,*
93 *Maryland*
94
95 Roberto Bilbao
96 *Barcelona Supercomputing Center, Barcelona, Spain*
97
98 Felipe M. de Andrade
99 *National Centre for Atmospheric Science, Department of Meteorology, University of Reading,*
100 *Reading, United Kingdom*
101
102 Michael J. DeFlorio
103 *Center for Western Weather and Water Extremes, Scripps Institution of Oceanography,*
104 *University of California, San Diego, California*
105
106 Leandro B. Díaz
107 *Universidad de Buenos Aires, Centro de Investigaciones del Mar y la Atmósfera, Institut Franco-*
108 *Argentin d’Estudes sur le Climat et ses Impacts, Buenos Aires, Argentina*
109
110 Muhammad Azhar Ehsan
111 *Earth System Physics Section, International Centre for Theoretical Physics, Trieste, Italy, and*
112 *Center of Excellence for Climate Change Research, King Abdulaziz University, Jeddah, Saudi*
113 *Arabia*
114
115 Georgios Fragkoulidis
116 *Institute for Atmospheric Physics, Johannes Gutenberg University, Mainz, Germany*
117
118 Sam Grainger
119 *Sustainability Research Institute, School of Earth and Environment, University of Leeds, Leeds,*
120 *United Kingdom*
121
122 Benjamin W. Green
123 *Cooperative Institute for Research in Environmental Sciences, University of Colorado, and*
124 *NOAA/OAR/ESRL/Global Systems Division, Boulder, Colorado*
125
126 Momme C. Hell
127 *Scripps Institution of Oceanography, La Jolla, California*
128
129 Johnna M. Infanti
130 *Cherokee Nation Strategic Programs, and NOAA/Office of Oceanic and Atmospheric*
131 *Research/Office of Weather and Air Quality, Silver Spring, Maryland*

132
 133 Katharina Isensee
 134 *Deutscher Wetterdienst, Offenbach, Germany*
 135
 136 Takahito Kataoka
 137 *Japan Agency for Marine-Earth Science and Technology, Kanagawa, Japan*
 138
 139 Ben P. Kirtman
 140 *University of Miami, Rosenstiel School for Marine and Atmospheric Sciences, Miami, Florida*
 141
 142 Nicholas P. Klingaman
 143 *National Centre for Atmospheric Science, Department of Meteorology, University of Reading,*
 144 *Reading, United Kingdom*
 145
 146 June-Yi Lee
 147 *Research Center for Climate Sciences, Pusan National University and Center for Climate Physics,*
 148 *Institute for Basic Science, Busan, Korea*
 149
 150 Kirsten Mayer
 151 *Department of Atmospheric Science, Colorado State University, Fort Collins, Colorado*
 152
 153 Roseanna McKay,
 154 *School of Earth, Atmosphere and Environment, Monash University, Melbourne, Australia*
 155
 156 Jennifer V Mecking
 157 *Ocean and Earth Science, University of Southampton, Southampton, United Kingdom*
 158
 159 Douglas E. Miller
 160 *University of Illinois at Urbana–Champaign, Urbana, Illinois*
 161
 162 Nele Neddermann
 163 *Institute for Oceanography, CEN, Universität Hamburg, and International Max Planck Research*
 164 *School on Earth System Modelling, Max Planck Institute for Meteorology, Hamburg, Germany*
 165
 166 Ching Ho Justin Ng
 167 *Atmospheric and Oceanic Sciences (AOS), Princeton University, Princeton, New Jersey*
 168
 169 Albert Ossó
 170 *NCAS-Climate, University of Reading, Reading, United Kingdom*
 171
 172 Klaus Pankatz
 173 *Deutscher Wetterdienst, Offenbach, Germany and Max Planck Institut für Meteorologie,*
 174 *Hamburg, Germany*
 175

176 Simon Peatman
 177 *School of Earth and Environment, University of Leeds, Leeds, United Kingdom*
 178
 179 Kathy Pegion
 180 *George Mason University, Fairfax, Virginia*
 181
 182 Judith Perlwitz
 183 *CIRES, University of Colorado Boulder, and Physical Sciences Division, NOAA/ESRL, Boulder,*
 184 *Colorado*
 185
 186 G. Cristina Recalde-Coronel
 187 *Department of Earth and Planetary Sciences, Johns Hopkins University, Baltimore, Maryland,*
 188 *and Facultad de Ingeniería Marítima y Ciencias del Mar, Escuela Superior Politécnica del Litoral,*
 189 *Guayaquil, Ecuador*
 190
 191 Annika Reintges
 192 *GEOMAR Helmholtz Centre for Ocean Research, Kiel, Germany*
 193
 194 Christoph Renkl
 195 *Department of Oceanography, Dalhousie University, Halifax, Canada*
 196
 197 Balakrishnan Solaraju-Murali
 198 *Barcelona Supercomputing Center, Barcelona, Spain*
 199
 200 Aaron Spring
 201 *Max Planck Institute for Meteorology, Hamburg, Germany*
 202
 203 Cristiana Stan
 204 *Department of Atmospheric, Oceanic and Earth Sciences, George Mason University, Fairfax,*
 205 *Virginia*
 206
 207 Y. Qiang Sun
 208 *Atmospheric and Oceanic Sciences Program, Princeton University, Princeton, New Jersey*
 209
 210 Carly R. Tozer
 211 *CSIRO Oceans and Atmosphere, Hobart, Tasmania, Australia*
 212
 213 Nicolas Vigaud
 214 *International Research Institute for Climate and Society (IRI), Columbia University, Palisades, NY*
 215
 216 Steven Woolnough
 217 *National Centre for Atmospheric Science, University of Reading, Reading, UK*
 218
 219 Stephen Yeager

National Center for Atmospheric Research, Boulder, Colorado, USA

Corresponding author: William Merryfield (bill.merryfield@canada.ca)

Abstract

Weather and climate variations on subseasonal to decadal timescales can have enormous social, economic and environmental impacts, making skillful predictions on these timescales a valuable tool for decision makers. As such, there is a growing interest in the scientific, operational, and applications communities in developing forecasts to improve our foreknowledge of extreme events. On subseasonal to seasonal (S2S) timescales, these include high-impact meteorological events such as tropical cyclones, extratropical storms, floods, droughts, and heat and cold waves. On seasonal to decadal (S2D) timescales, while the focus broadly remains similar, (e.g., on precipitation, surface and upper ocean temperatures and their effects on the probabilities of high-impact meteorological events), understanding the roles of internal and externally-forced variability such as anthropogenic warming in forecasts also becomes important.

The S2S and S2D communities share common scientific and technical challenges. These include forecast initialization and ensemble generation; initialization shock and drift; understanding the onset of model systematic errors; bias correction, calibration, and forecast quality assessment; model resolution; atmosphere-ocean coupling; sources and expectations for predictability; and

linking research, operational forecasting, and end user needs. In September 2018 a coordinated pair of international conferences, framed by the above challenges, was organized jointly by the World Climate Research Programme (WCRP) and the World Weather Research Programme (WWRP). These conferences surveyed the state of S2S and S2D prediction, ongoing research, and future needs, providing an ideal basis for synthesizing current and emerging developments in these areas that promise to enhance future operational services. This article provides such a synthesis.

Capsule

Climate prediction on subseasonal to decadal time scales is a rapidly advancing field that is synthesizing improvements in climate process understanding and modeling to improve and expand operational services worldwide.

[Introductory text]

Beyond the tremendous progress in weather forecasting witnessed in recent decades (Bauer et al. 2015), predictive capabilities have expanded, increasingly seamlessly, to encompass climate on subseasonal to decadal time scales (Fig. 1 and Kirtman et al. 2013). These advances have been enabled by better observations, data assimilation schemes, and models originating both from the weather prediction and long term climate simulation communities, together with increased computational power supporting progressively higher resolution and larger ensembles that allow uncertainties to be better estimated and in some cases reduced.

International efforts under the auspices of the World Weather Research Programme (WWRP) and World Climate Research Programme (WCRP) have helped drive this progress through coordinated research to improve the accuracy and utilization of weather and climate predictions. Community research efforts under the WCRP led initially to climate predictions one to two seasons ahead becoming part of the World Meteorological Organization (WMO) operational infrastructure (Graham et al. 2011). More recently a joint WWRP and WCRP Subseasonal to Seasonal Prediction Project has started tackling the so-called weather-climate prediction desert from two weeks to a season (Robertson et al. 2018; Mariotti et al. 2018), aiming to underpin new WMO operations on those time scales (Vitart et al. 2017), and the NOAA-led SubX project has similar aims (Pegion et al. 2019). At longer ranges, WCRP-enabled research has quantified predictability from a year to a decade, and corresponding WMO operational infrastructure for annual-to-decadal climate prediction is now in place (World Meteorological Organization 2018; Kushnir et al. 2019).

292
293 As each of these efforts has progressed it has become increasingly apparent that common
294 challenges exist across predictive time scales. These include understanding and adequately
295 representing in models processes that give rise to predictability in the Earth system, consisting
296 of the physical climate system—atmosphere, ocean, land and sea ice—together with associated
297 biogeochemical cycling, especially of carbon (upper part of Fig. 1); capturing and
298 communicating inherent uncertainties caused by the chaotic nature of weather and climate;
299 correcting for and reducing imperfections in models that may systematically degrade forecast
300 quality; and providing forecast information in a form that is applicable to decision making. At
301 the same time, opportunities for usefully predicting elements of the Earth system beyond long-
302 term means of standard meteorological variables , including land, ocean and sea ice properties
303 and risks of weather extremes, have come into focus. The ultimate collective endeavor is to
304 improve the prediction of the spatial–temporal continuum connecting weather to climate
305 through a coordinated, seamless and integrated Earth system approach for the benefit of
306 society.

307
308 In September 2018, international conferences¹ on subseasonal to seasonal prediction (S2S,
309 encompassing forecast ranges from two weeks to a season) and seasonal to decadal prediction (S2D,
310 encompassing ranges longer than a season, up to a decade) together with cross-cutting plenary

¹ The Second International Conference on Subseasonal to Seasonal Prediction (S2S) and Second International Conference on Seasonal to Decadal Prediction (S2D) were held 17-21 September 2018 at NCAR facilities in Boulder Colorado. These coordinated meetings involved 347 participants, including 92 early career scientists, from 38 countries, with a total of 368 oral and poster presentations. Further information including a complete list of contributions can be found at <https://www.wcrp-climate.org/s2s-s2d-2018-home>.

311 sessions were convened jointly by WWRP and WCRP. This represented a confluence of research
312 and operational climate prediction expertise and knowledge exchange across prediction time
313 scales that was unprecedented in scope. Selected outcomes, organized by themes
314 encompassing the challenges outlined above, are synthesized in this article.

316 **Mechanisms of predictability.**

317 *Subseasonal to Seasonal*

318 A major source of S2S predictability is the organization of tropical convection by the Madden
319 Julian Oscillation, or MJO (Woolnough, 2019), which is predicted skillfully by S2S project models
320 up to 3-4 weeks ahead (Vitart 2017). The MJO has worldwide impacts that depend on its
321 amplitude and phase, including modulation of tropical cyclone activity (Lee et al. 2018; Zhao et
322 al. 2019) and extratropical phenomena such as the East Asian summer monsoon (Li et al. 2018).
323 The associated tropical-extratropical teleconnections (Lin et al. 2019) impart S2S forecast skill
324 for many of these extratropical phenomena including Euro-Atlantic weather regimes, position
325 of the jet stream, atmospheric rivers (DeFlorio et al. 2019), and hail/tornado activity (Baggett et
326 al. 2018). However, good representations of the basic state both in the tropics and extratropics,
327 as well as tropical air-sea interactions and atmospheric convection (e.g., Yoo et al. 2015), are
328 necessary for these teleconnections to be correctly simulated by general circulation models
329 (Henderson et al. 2017).

S2S predictability also derives from the stratosphere through its relatively long time scales of variability² and lagged influences on the troposphere (Kidston et al. 2015). Interactions between the stratosphere and the troposphere from the tropics to the extratropics thus provide a promising source of S2S prediction skill (Butler et al. 2019). For example, in the winter Northern Hemisphere stratosphere the climatological westerly polar vortex exhibits extremes in variability, including sudden stratospheric warmings (SSWs) that are driven largely by Rossby waves from the troposphere. SSWs have lagged impacts on sea level pressure, surface temperature and precipitation, including pronounced tendencies for cold anomalies over northern Eurasia and warm anomalies over northeastern North America (e.g., Sigmond et al. 2013). Initializing forecasts during extreme stratospheric events provides increases in prediction skill of surface climate in such regions up to 3-6 weeks later (Domeisen et al. 2019c). However, the predictability of specific extreme stratospheric events is limited, ranging from a few days to about two weeks (Fig. 2) for different SSWs (Karpechko 2018; Taguchi 2018, Domeisen et al. 2019a), although models show evidence of under-confident forecasts in the stratosphere on S2S timescales (O'Reilly et al. 2019). Outstanding questions remain about the mechanisms of stratosphere-troposphere coupling processes, in particular on the causes, variability, and trends for the occurrence of SSW events (Ayarzaguena et al. 2018; Simpson et al. 2018) and why not all SSW events have similar downward effects (e.g., Garfinkel et al. 2013, Maycock & Hitchcock, 2015). In addition, further research is needed to assess the degree to which prediction models capture both the stratospheric variability and coupling processes.

² Including the quasi-biennial oscillation (QBO) of the tropical stratosphere, whose influences span a range of time scales and are addressed in the “Time scale interactions” subsection.

351
352 Among atmosphere-surface influences, land-atmosphere interactions have their greatest
353 impact on subseasonal time scales in forecasts where land is initialized (Dirmeyer et al. 2018a),
354 but also can contribute skill on weather prediction and multi-month time scales (Dirmeyer and
355 Halder 2016, 2017). The most broadly impactful land attribute is soil moisture (Koster et al.
356 2004, 2016), but anomalies in soil temperature (Y. Zhang et al. 2019; Yang et al. 2019), snow
357 cover (Jeong et al. 2012; Orsolini et al. 2013), and vegetation states (Williams et al. 2016) can all
358 have significant impacts. A number of recent studies have focused on non-local impacts of land
359 surface anomalies, showing for example that soil moisture anomalies can exert remote as well
360 as local influences in boreal summer through driving of quasi-stationary Rossby waves and
361 associated circulation anomalies (e.g., Teng et al. 2019; Wang et al. 2019). In addition, land
362 surface and subsurface temperatures in spring may exert delayed downstream influences on
363 precipitation (Xue et al. 2018), and evapotranspiration may remotely influence precipitation
364 over land (Wei and Dirmeyer 2019).

365
366 Atmosphere-ocean interactions, fundamental for S2D predictability, can also be influential on
367 S2S time scales. For example submonthly prediction skills for precipitation and temperature are
368 enhanced over certain land areas including parts of Australia, the Maritime Continent and the
369 contiguous United States when tropical sea surface temperature (SST) anomalies associated
370 with El Niño Southern Oscillation (ENSO) are present (Hudson et al. 2011; Li and Robertson
371 2015; DelSole et al. 2017). Extratropical SST anomalies also can impart S2S skill through
372 teleconnections, as shown for example by McKinnon et al. (2016) who identified a SST anomaly

pattern in the mid-latitude North Pacific that tends to precede heat waves and rainfall deficits in the eastern United States by up to 50 days.

Sea ice strongly influences surface fluxes and lower atmospheric temperatures particularly in the marginal ice zone, and provides a source of S2S predictability for polar and possibly midlatitude regions (Chevallier et al. 2019). This motivates the development of S2S forecasts for sea ice, which thus far have shown significant, albeit region-dependent skill for predicting intraseasonal Arctic sea ice variability (Liu et al. 2018, Zampieri et al. 2018).

Seasonal to decadal

A primary general source of S2D atmospheric predictability is remote influences from a variety of teleconnections (e.g., Yuan et al. 2018; Ruprich-Robert et al. 2018; Beverley et al. 2019). Teleconnections associated with anomalous atmospheric circulation patterns arise from changes to the Walker circulation usually driven by anomalous zonal SST gradients (Cai et al. 2019), and changes to the Hadley circulation usually driven by anomalous meridional SST gradients, especially interhemispheric differences (Kang et al. 2018). These influences impact tropical cyclones and rainfall, whereas anomalous upper level divergence due to tropical rainfall anomalies leads to Rossby waves that impact the extratropics (Scaife et al. 2017; O'Reilly et al. 2018). Besides giving rise to atmosphere-ocean interactions that alter the atmospheric circulation, SST anomalies can induce low-level temperature and moisture anomalies that are advected elsewhere by climatological winds (Dunstone et al. 2018).

S2D atmospheric predictability arising from teleconnections requires that SST anomalies be predictable. On seasonal timescales, tropical SST anomalies are dominated by ENSO (Yang et al. 2018), though there is some independent variability in the tropical Atlantic and Indian Oceans that also drives teleconnections (e.g., Nnamchi et al. 2015; Lim et al. 2016). The impacts of ENSO are sensitive to ENSO diversity (Capotondi et al. 2015), including the longitude at which maximum SST anomalies occur (Yeh et al. 2018; Patricola et al. 2018). ENSO SST anomalies are largely predictable out to a year particularly in winter and early spring (Barnston et al. 2017), whereas predictability may extend to two years for some La Niña events (Di Nezio et al. 2017), and to 1 ½ to two years for certain El Niño events (Luo et al. 2008).

Decadal SST variability occurs in both the Atlantic and Pacific oceans, often referred to as Atlantic Multidecadal Variability (AMV) and Pacific Decadal Variability (PDV), e.g. Kushnir et al. (2019). The causes of AMV are not fully understood, especially the relative roles of internal variability and external forcing from aerosols. However, AMV is modulated to some extent by the oceanic Atlantic Meridional Overturning Circulation (Yeager and Robson 2017), which together with the North Atlantic subpolar gyre is influenced by deep ocean density anomalies particularly in the Labrador Sea (Robson et al. 2016); these influences contribute to the especially high multi-year predictability in the North Atlantic (Buckley et al. 2019). AMV couples to the Hadley circulation, affecting hurricanes and Sahel rainfall as illustrated in Fig. 3 (Sheen et al. 2017), and can initiate atmospheric Rossby waves with remote influences including temperatures in parts of China (Monerie et al. 2018). AMV can influence PDV (Ruprich-Robert et al. 2017), and vice-versa. PDV may also be influenced by off-equatorial heat content

anomalies in the western Pacific Ocean (Meehl et al. 2016). Decadal variability of deep convection in the Southern Ocean influences temperatures in that region, potentially explaining recent increases in Antarctic sea ice (L. Zhang et al. 2019).

S2D atmospheric predictability also arises from longer time scale processes over land, mainly involving soil moisture (Chikamoto et al. 2017; Ardilouze et al. 2019) and vegetation (Weiss et al. 2014; Bellucci et al. 2015). These highlight the need for land surface initialization (Prodhomme et al. 2016a) and realistic vegetation models (Alessandri et al. 2017).

An additional source of S2D predictability is variations in radiative forcing, which provide significant skill on multi-year timescales (Smith et al. 2019). Much of this skill arises from changes in greenhouse gases, but anthropogenic aerosols may force decadal variations in AMV (Booth et al. 2012) and PDV (Smith et al. 2016; Takahashi and Watanabe 2016). Solar variability (Misios et al. 2019), and volcanic eruptions (Menegoz et al. 2018) including through their influence on ENSO (Khodri et al. 2017; Wang et al. 2018) and possibly AMV and the North Atlantic Oscillation (NAO; Swingedouw et al. 2017) affect climate on seasonal to decadal timescales and are potentially important sources of predictability. However, the relative roles of external radiative forcing and internal variability (W. Kim et al. 2018) continue to be explored.

Time scale interactions

438 The Quasi-biennial Oscillation (QBO) is a downward-propagating ~28-month oscillation of
439 easterly and westerly zonal jets in the tropical stratosphere, driven by upward equatorial waves
440 from the troposphere (e.g., Kim and Chun 2015). In addition to having high predictability and
441 some teleconnected influence on winter surface climate (e.g., Scaife et al. 2014a), the QBO
442 modulates the amplitude, persistence, and rate of propagation of the boreal wintertime MJO
443 (Fig. 4) through its impact on tropical convection via changes in static stability near the
444 tropopause (Yoo and Son 2016, Nishimoto and Yoden 2017). MJO amplitude is better predicted
445 at longer leads during the easterly phase of the QBO (Marshall et al. 2017), likely as a result of
446 longer persistence of the MJO rather than its greater initial amplitude (Lim et al. 2019).

447

448 The modulation of SSW probability of occurrence by tropical sources of variability, such as the
449 QBO, ENSO, or MJO, may extend probabilistic predictability of stratospheric variability to a few
450 months or longer if these relationships can be adequately captured by prediction models
451 (Garfinkel & Schwartz 2017; Garfinkel et al. 2018; Domeisen et al. 2019a,b).

452

453 There is increasing evidence of additional interactions between various sources of S2S and S2D
454 predictability across time scales. One example is that seasonal time scale variations in ENSO
455 modulate the MJO (Chen et al. 2016) and its impact on the NAO (Lee et al. 2019) with
456 consequent influences on weather over remote regions. Another is that ENSO teleconnection
457 to the extratropics has varied over multi-decadal time scales spanning the past 100+ years
458 (O'Reilly 2018), possibly modulating ability to predict the NAO (Weishiemer et al. 2019),

although sampling variability can also give rise to such long-term changes in teleconnections (Yun and Timmermann 2018).

Modelling issues.

Subseasonal to Seasonal

Because S2S operational prediction is a relatively new enterprise, considerable efforts focusing on fundamental aspects of forecast system design are occurring at operational centers worldwide (Takaya, 2019). One major emphasis consists of methods to represent the uncertainty in initial conditions (bred vector, singular vector, ensemble data assimilation) and model physics (stochastic physics, Leutbecher et al. 2018). In addition, configurations of real-time forecasts and hindcasts, including ensemble size, ensemble strategy (lagged ensemble with different initial times or burst ensemble with a common initial time) and hindcast period, impact forecast quality and ability to evaluate the performance of the hindcast. Identifying suitable compromises and trade-offs in forecast system design is a challenge under practical constraints for operational activities (costs, priorities, timeliness) and demands further research.

From the modelling perspective, multiple operational centers are moving towards a unified, or “seamless” coupled forecast system that can be applied across timescales from days to seasons or longer. More S2S models are incorporating ocean and sea-ice components, and becoming increasingly complex and complete in representing coupled processes in the Earth system. On the other hand, poor representation of model physics, in particular clouds (Morcrette et al.

2018), results in model drifts and biases in surface land and ocean temperatures, which is a long-standing modeling issue that can degrade the skill of S2S predictions (Vitart and Balmaseda, 2017). Improvements in cloud parameterizations (Stan and Straus, 2019) and in representing the diurnal cycle of the atmospheric boundary layers are crucial for advancing S2S modeling. The Earth system modeling approach poses another challenge to initialize the ocean and sea ice components with high accuracy; for example there is a relatively large dispersion of initialized sea ice fields in current S2S models (Chevallier et al. 2017, Zampieri et al. 2018). Another important S2S modeling issue is predicting the MJO, owing to its importance as a source of subseasonal predictability (H. Kim et al. 2018). Multi-model evaluations have shown that S2S models have difficulties in representing MJO propagation across the Maritime Continent. Process-oriented diagnostics (Maloney et al. 2019) have identified a dry bias in the lower troposphere as one of the causes for the poor MJO propagation through weakening the horizontal moisture gradient over the Indian Ocean and western Pacific (Lim et al. 2018) and dampening the organization and propagation of the MJO. A recharge process whereby moisture builds up in the lower troposphere during the suppressed convection phase of the MJO, and that is key for MJO propagation around the Maritime Continent in boreal winter, is underrepresented in S2S models due to the dry bias (Kim 2017). Ocean coupling is another important process for the MJO (DeMott et al. 2015), and several studies have demonstrated that ocean coupling can improve MJO propagation and enhance predictive skill in models.

Poor vertical resolution, low model lid height, inadequate orographic and non-orographic gravity wave parameterizations, and biases in the tropospheric mean state (e.g., the location of

stationary Rossby waves) could limit the predictive skill from stratosphere-troposphere coupling processes (Tripathi et al. 2015; Butler et al. 2016), but new generations of prediction systems have rapidly improved in many of these areas. Future model development could prioritize improved representation of orographic and non-orographic gravity wave drag and an internally-generated QBO (Butchart et al. 2018). Better understanding of stratosphere-troposphere coupling processes and the role of the stratosphere on surface skill could be gained through case studies and stratospheric nudging experiments (Hansen et al. 2017). Improved observations of the stratosphere (e.g., aerosols and chemistry) and other climate components may improve S2S predictions. Finally, there is potential for modeling of stratospheric ozone chemistry which provides surface temperature predictability on S2S time scales due to its influence on high-latitude stratospheric circulation anomalies together with their lagged surface impacts (Stone et al. 2019). Although that may currently be too resource-intensive due to the many species and reactions that must be modeled, emerging machine-learning techniques may provide pathways for incorporating chemistry-climate information into S2S forecasts (Nowack et al. 2018).

Seasonal to decadal

Modeling issues for S2D prediction naturally overlap with those for S2S prediction. However, the longer time scales of S2D prediction lead to a greater emphasis on representing slower climate variations such as ENSO and AMV, and greater attention to reducing model biases in the ocean that may take months to years to develop. Increased model resolution can reduce model biases as illustrated in Fig. 5 (Jia et al. 2015; Müller et al. 2018), and improve skill

(Prodhomme et al. 2016b; Schuster et al. 2019; Infanti and Kirtman 2019), although the greater computational cost is not always justified (Scaife et al. 2019). More fundamental strategies involve analyzing/understanding model biases, before attempting to correct them a priori or a posteriori. Such analysis methods include comparing hindcasts with observations and multi-decadal historical or other simulations to distill causation for model errors, such as in the tropical Pacific (Shonk et al. 2018) or Atlantic (Voldoire et al. 2019). Similarly, errors in modeled variability or teleconnection patterns can be characterized by examining their evolution with lead time. Model biases can be corrected both through simple methods such as statistical bias correction and anomaly coupling (Toniazzo and Koseki, 2018), and more complex methods such as supermodeling, through which multiple models exchange information during a climate simulation (Shen et al. 2016).

Performance of S2D predictions is strongly tied to initialization of model components beyond the lower atmosphere. For example, stratospheric initial conditions are a source of seasonal winter NAO skill (e.g., O'Reilly et al. 2019; Nie et al. 2019) as illustrated in Fig. 6, and ocean initial conditions are crucial for skillfully predicting ENSO (Balmaseda and Anderson 2009), as well as decadal variability in the subpolar North Atlantic (Yeager and Robson, 2017; Borchert et al. 2018). However, initialization using full-field observational values can lead to initial shocks affecting skill (Kröger et al. 2018) and in such cases initialization combining observed anomalies with the model's own climatology can be beneficial until underlying model errors can be reduced (Volpi et al. 2017). Basic initialization strategies continue to be an active research area particularly for decadal prediction (Brune et al. 2018), and methods extending to forecast runs

such as the ensemble dispersion filter which replaces the ensemble members with the ensemble mean every three months (Kadlow et al. 2017) are also being explored. Comparisons that apply different initialization methods to the same model can yield valuable insights (Polkova et al. 2019); further issues specific to the initialization of the land, ocean, and sea ice components are considered in the next section.

Tackling these diverse and persistent modeling issues effectively will require sustained effort, as simple model-specific solutions may not cure the underlying problems, and ideally this should involve coordination between the S2S/S2D prediction, climate modelling, and data assimilation communities.

Initialization issues.

Atmosphere initialization

Accurate atmospheric model initialization is a basic requirement for numerical weather prediction because atmospheric initial conditions are the primary source of predictability on time scales less than a week or two (Fig. 1). It is enabled by sophisticated data assimilation systems that are the result of decades of advancement (Bauer et al. 2015). Subseasonal and seasonal prediction systems generally initialize their atmospheric components by such means, with the additional requirement that historical observations must be assimilated similarly to produce reanalyses that are used to initialize hindcasts. Because in situ and remotely sensed atmospheric observations are relatively dense there is generally good agreement between different reanalyses for the modern era implying relatively low uncertainty at heights below

about 10 hPa, although temporal inconsistencies can result from changes in observing systems (Long et al. 2017). Because atmospheric initial conditions contribute less to predictability on multi-annual time scales, some decadal prediction systems do not initialize the atmosphere (e.g., Yeager et al. 2018).

Land initialization

Climatically important land variables such as soil moisture and snow can be initialized by driving land surface models with observed atmospheric fields (e.g., Koster et al. 2009; Sospedra-Alfonso et al. 2016a) or, more directly, assimilation of land observations principally from satellites (Bilodeau et al. 2016; Muñoz-Sabater et al. 2019; Toure et al. 2018). Yet predictability from land surface states is being harvested only to the extent that land initial conditions and the relevant processes are represented realistically in forecast models (Koster et al. 2011; Ardilouze et al. 2017). Historically, land surface and atmospheric models are developed separately and their coupled behavior is not calibrated or validated (Dirmeyer et al. 2019), so that coupled processes are often not represented accurately (Dirmeyer et al. 2018b).

There are also observational limitations. In situ measurements of soil moisture are of varying quality and uneven distribution, and are not designed for real-time operational use (Dorigo et al. 2011). Satellite soil moisture monitoring (Entekhabi et al. 2010; Kerr et al. 2010), provides either very shallow or total column measurements including groundwater (Li et al. 2012), and is subject to uncertainties caused by vegetation, etc. (Al-Yaari et al. 2017). By contrast, soil moisture in forecast models is mainly a gross reservoir for the surface water balance, and its

variations do not represent all of the observed processes, particularly at sub-grid scales. Therefore model soil moisture is only a crude representation of reality, although it still contains useful information that can be largely consistent across different land models (Koster et al. 2009).

Climate forecasts can be improved by making high-quality land state observations an operational priority for real-time reporting, and planning for long-term continuity in satellite monitoring (Balsamo et al. 2018). This includes vegetation, especially as its interannual variability and cycles of agricultural planting and harvest are not represented and can affect surface fluxes and predictions (Alessandri et al. 2017). In addition, realistic snow initialization can positively impact subseasonal predictions of surface temperatures (e.g., F. Li et al. 2019). Along with coupled land-atmosphere model development (Santanello et al. 2018), such efforts would facilitate improved predictions on weather to subseasonal time scales, as demonstrated by numerous forecast model-based sensitivity studies such as that of Koster et al. (2011).

Ocean and sea ice initialization

The importance of initializing the oceans stems from their relatively long thermal and dynamical time scales, through which they play an essential role in S2D climate predictability (Cassou et al. 2017). In addition, the oceans can influence S2S variability, for example through air-sea interactions affecting the MJO (DeMott et al. 2015) and mesoscale eddy impacts on atmospheric circulation (Saravanan and Chang 2019). Predicting future ocean evolution, especially on S2D time scales, requires estimates of 3D ocean states for initialization. This in

turn requires a data assimilation method (usually in conjunction with a dynamical model) to constrain ocean state estimates based on available observations. Similar considerations apply to state estimates of sea ice. Comparisons of different ocean and sea ice state estimates as in Fig. 7 can point to variables and regions for which they are most robust, as well as to where uncertainties are relatively large (Balmaseda et al. 2015; Chevallier et al. 2017). Observing system experiments in which certain observations are withheld have shown for example that data from tropical ocean moorings positively impacts state estimates even when Argo float data is also available (Fujii et al. 2015).

Recent enhancements in observing capabilities are enabling improvements in ocean and sea ice state estimates, potentially leading to more accurate initial conditions and hence better forecasts. For example, assimilation of satellite measurements of sea surface salinity (SSS) leads to improvements in tropical Pacific ocean states and ENSO forecasts in experiments using an intermediate-complexity coupled model (Hackert et al. 2019), whereas assimilation of satellite-derived sea ice thickness (SIT) measurements has shown potential for improving sea ice forecasts in operational seasonal forecasting systems (Chen et al. 2017; Blockley and Peterson, 2018). A major limitation is that these data sources have been available for less than a decade, whereas considerably longer hindcast periods are required for forecast post-processing and skill assessment, and temporal consistency of observational data used for initialization is required to avoid artificial biases between hindcasts and forecasts. Forecasts thus continue to be initialized typically without assimilation of SSS or SIT, from initial conditions that deviate appreciably from

available observations especially for SIT (Uotila et al. 2019). This motivates alternative approaches for initializing SIT over multidecadal hindcast periods (Dirkson et al. 2017).

Coupled data assimilation

The atmosphere, land, ocean and sea ice components of climate prediction models have often been initialized individually, without coupling. However, such an approach does not make optimal use of observations, which may exert influences across the interfaces of the model components. In addition, physical inconsistencies between the separately initialized components may lead to rapid adjustments, or shocks. To overcome these limitations attention has increasingly turned toward developing coupled data assimilation methods that treat multiple components, such as atmosphere and ocean, simultaneously using observations from each (Penny and Hamill 2017). Such methods are termed weakly or strongly coupled (Penny et al. 2017). Weakly coupled methods apply assimilation independently to each model component within the coupled model, so that the components may exchange information across their interfaces. Such techniques have shown promise for reducing shocks (Mulholland et al. 2015), and have begun to be applied operationally (e.g., Browne et al. 2019). Strongly coupled methods apply assimilation to multiple model components in an integrated manner, so that observations assimilated in one component can directly influence others. Such methods remain experimental and thus far have been applied mainly in simplified models (e.g., Penny et al. 2019).

Ensemble predictions and forecast information.

656 *Subseasonal to Seasonal*

657 In contrast to ensemble weather forecasts, a consolidated verification strategy for S2S
658 predictions is not yet established, and developing such a framework that encompasses
659 important forecast attributes such as accuracy, association, discrimination, reliability, and
660 resolution has thus emerged as a priority (Coelho et al. 2018). (Accuracy measures error, or
661 distance between forecast and observed values; association measures strength of the linear
662 relationship between forecast and observation as through temporal or spatial correlations;
663 discrimination measures by how much forecasts differ given different outcomes; reliability
664 measures how well forecast probabilities correspond to observed frequencies of occurrence;
665 resolution measures by how much outcomes differ given different forecast probabilities.
666 Forecast quality encompasses all these attributes, whereas skill indicates quality relative to
667 some benchmark such as persisted anomalies or climatological probabilities.) As for seasonal
668 predictions, a purpose of S2S hindcasts is to provide a larger sample for more confident
669 verification statistics than real time forecasts because they cover more years. However,
670 because S2S hindcasts are initialized using re-analysis and most often have a smaller ensemble
671 size, their verification generally underestimates real-time forecast quality. Operational centres
672 are encouraged to compute and monitor verification statistics based both on hindcasts and
673 real-time forecasts.

674

675 As has been demonstrated for seasonal prediction, S2S multi-model ensembles (MMEs)
676 generally outperform individual models (Vigaud et al. 2017; Pegion et al. 2019). Currently, the

S2S and SubX MME projects are providing testbeds for research³ as well as a foundation for operational use (Vitart and Robertson 2019; Pegion et al. 2019). One focus for exploiting such datasets is developing calibration procedures, post-processing steps that improve the properties of probabilistic forecasts, to enable S2S ensemble forecasts to provide reliable probabilities for particular conditions occurring or thresholds being exceeded, especially for extreme events. The varied current choices among S2S project modelling systems for hindcast and near real time initialization dates, hindcast period and ensemble size is, however, limiting advances in developing multi-model calibration and combination procedures. In addition, the value of these datasets for research would be enhanced if more comprehensive stratospheric data were to be available across models.

S2S ensemble forecasts have shown promise in providing useful predictions and early warnings for high impact climate and weather events including severe heat waves and cold spells, as well as regional probabilities of the occurrence of tropical storms as illustrated in Fig. 8 (Vitart and Robertson 2018). Examples include severe cold conditions over Europe associated with the negative phase of the NAO, whose useful predictability into week 3 is enhanced by tropical–extratropical teleconnections resulting from MJO activity (Ferranti et al. 2018), and atmospheric rivers, plumes of intense water vapor transport that often trigger weather and hydrologic extremes and are especially predictable at lead times of 3 to 5 weeks during certain MJO and QBO phase combinations (Baggett et al. 2017). While modest overall skill at ranges longer than

³ Hindcast and near real-time forecast data are available from S2S at www.s2sprediction.net and from SubX at <http://iridl.ldeo.columbia.edu/SOURCES/.Models/.SubX/>.

a week has been found for S2S predictions of springtime Sahelian heat waves including measures of heat stress, such conditions following a strong El Nino were accurately forecast, pointing to the tropical Pacific as a source of predictability for extremes in that region (Batté et al. 2018).

A global precipitation hindcast quality assessment of the S2S prediction project models (Fig. 9) was performed by de Andrade et al. (2019). Sub-seasonal prediction quality is modulated by the MJO, QBO, ENSO in the tropics, changes in large-scale flow in the extra-tropics and stratospheric tropical and extratropical variability (Butler et al. 2019). It is therefore important to estimate the predictive skill of such events and identify their impacts on predictions of weather and weather extremes. Evaluating the conditional prediction quality associated with the key low frequency variability modes is instrumental for better understanding S2S predictability mechanisms. For example, MJO predictive skill in the S2S MME ranges between 12 to 36 days and is affected both by the MJO amplitude and phase errors (Vitart 2017; Lim et al. 2018; H. Kim et al. 2018). Communicating these variations in forecast quality, including if the forecasts are no better than climatology, is extremely important as users with such knowledge can better utilize and benefit from the forecast information. Furthermore, capitalizing on “windows of opportunity” when skill is especially high increases the value of S2S forecasts (Mariotti et al. 2020), and motivates their frequent initialization (ideally daily).

Seasonal to decadal

718 Limited forecast quality in current S2D ensemble prediction systems motivates research
719 initiatives that focus on extracting skillful and reliable information from the large amounts of
720 forecast and hindcast data available to potential users⁴.

721
722 One emerging theme of such research is that S2D prediction systems sometimes underestimate
723 the predictable signal (Eade et al. 2014; Scaife and Smith 2018). As a result, very large ensembles
724 that effectively filter out unpredictable noise demonstrate higher skill in predicting phenomena
725 such as the winter NAO (Scaife et al. 2014b; Dunstone et al. 2016) and seasonal to multi-annual
726 regional precipitation variations (Dunstone et al. 2018; Yeager et al. 2018) than was previously
727 thought possible. While very large ensemble sizes hold value for isolating weak predictable
728 signals, much smaller ensemble sizes are sufficient for skillful prediction of tropical SST, for which
729 signal to noise ratios are much larger (Zhu et al. 2015). The causes of unrealistically low modeled
730 predictable signals (sometimes called the “signal to noise paradox”) remain under investigation.
731 Two hypotheses stemming from hindcast experiments are that winter NAO skill is enhanced by
732 skillful prediction of a QBO teleconnection that is too weak in models (O’Reilly et al. 2019), and
733 that transient eddy feedbacks are too weak in models (Scaife et al. 2019). Others based on simple
734 models suggest that the NAO predictable signal is too weak because climate models switch

⁴ Seasonal hindcast data from the WCRP Climate-system Historical Forecast Project (CHFP; Tompkins et al. 2017) are available at <http://chfps.cima.fcen.uba.ar/access.php>, and from the North American Multi-Model Ensemble (NMME, Kirtman et al. 2014) including real-time forecasts at <https://iridl.ldeo.columbia.edu/SOURCES/.Models/.NMME/>. Decadal hindcast data from the WCRP Coupled Model Intercomparison Project Phases 5 and 6 are available via <https://esgf-node.llnl.gov/projects/cmip5/> and <https://esgf-node.llnl.gov/projects/cmip6/>.

between NAO regimes too rapidly (Strommen and Palmer 2019), or exhibit less persistent NAO variability than is observed (Zhang and Kirtman 2019).

In the case of the winter NAO which is a key source of variability over the mid-latitude North Atlantic and Europe, another approach to extract relevant information from over-dispersive ensembles that leads to improved skill is to subsample ensemble members that are close to a “first guess” statistical prediction of the NAO (Dobrynin et al. 2018); subsampling has shown potential for improving European summer forecasts as well (Neddermann et al. 2019).

Estimating and realizing the predictability of key modes of variability is still a major challenge at S2D time scales. ENSO is considered one of the most predictable phenomena on multi-seasonal time scales, but longer-range skill has been viewed as limited. However, multi-year ensemble predictions have shown evidence of skill in predicting long-lasting La Niña events that follow warm events up to 24 months ahead (DiNezio et al. 2017; Luo et al. 2017). Challenges in the initialization of such longer time scale predictions remain, as evidenced by multi-year predictions in which skill for SST and precipitation over land improves with lead time in some areas, suggesting that short-term adjustments following initialization may tend to degrade skill (Yeager et al. 2018).

Calibration of ensemble forecasts is a necessary step to reduce the areas for which S2D forecasts are unreliable and potentially misleading. Combinations of several forecasting systems such as the North American Multi-Model Ensemble (NMME, Kirtman et al. 2014) are now routinely used

to increase ensemble reliability and improve forecast skill. Several recent efforts have explored weighted multi-model calibration methods to combine ensembles from different models in order to improve probabilistic seasonal forecasts for temperature and precipitation anomalies as well as forecast of extremes (Becker 2017). Calibration methods have also been developed for ensemble decadal hindcasts to adjust both the bias and ensemble spread with a parametric dependency on lead time and initialization time (Pasternack et al. 2018). Such methods are found to improve both the conditional bias and probabilistic skill of decadal hindcasts.

Climate forecasts for decision making.

Subseasonal to Seasonal

Many decisions are made on the S2S forecasting timescale, which sits between weather forecasts and S2D climate outlooks; therefore the continued development of S2S forecasts has the potential to benefit many sectors of society (Fig. 10). S2S forecasting is a rapidly maturing discipline, with emerging recognition for both the need and the potential use of forecasts on this timescale (White et al. 2017). While S2S forecasts are increasingly being used in government as well as a range of sectors including agriculture, energy, finance, health and water resource management – more engagement between S2S forecasters and end users is needed to increase the wider awareness and uptake of S2S forecasts.

Although scientific knowledge gaps, computational capacity, and gaps in observations and modeling currently limit the use of S2S forecasts to some degree, by increasingly placing decision makers at the forefront of S2S forecast development, an improved dialogue between

779 S2S forecasters, developers and end users will accelerate the awareness and application of S2S
780 forecasts to real-world decision-making.

781

782 To support the increased use of S2S forecasts for decision-making, the following
783 recommendations were identified for action following the Boulder conference:

- 784 ● A summary of existing stakeholder case studies is planned to be created to demonstrate past
785 and ongoing ‘success stories’, and support better engagement with end users and
786 stakeholders. As the S2S forecast needs and associated performance varies greatly between
787 different sectors and users, the wider community is increasingly working together on the co-
788 design and production of S2S predictions in order to better meet user needs. Several
789 applications of S2S forecasts are now being developed in different disciplines, such as the
790 EU-funded S2S4E project in the energy sector, a quasi-operational excess heat outlook
791 system in the health sector (Lowe et al. 2016), and S2S hydrologic prediction in the water
792 management sector. These efforts need to be catalogued and disseminated to guide further
793 user-driven decision-support products, and to support the continued development of S2S
794 forecast, verification metrics and related services.
- 795 ● Systematically assessing the relative skill (or lack thereof) of forecasting a series of historical
796 high-impact events, such as heat waves, extreme rainfall events, or wildfires, on the S2S
797 timescale would be a useful way to help demonstrate the potential of S2S forecasts to
798 decision-makers across multiple sectors. At present, such case studies are often ad-hoc and
799 typically not published in the wider literature; however, a collaborative effort that brings
800 together a set of demonstrable case studies, involving both forecasters and end users, would

fill this gap. For example, a series of ‘tailored narratives’, or ‘storylines’ (approaches that construct stories of plausible, non-probabilistic climatic futures that relate to a specific person or sector to counter perceived barriers; e.g., Hazeleger *et al.* 2015), may aid in the understanding of what S2S forecasts may deliver in the future.

- To support the co-design, uptake and use of S2S forecasts, S2Sapp.net is currently being established as a new network of researchers, modellers and practitioners – an ‘open to all’ global community with a shared aim of exploring and promoting cross-sectoral services and applications of this new generation of forecasts from across government, academia, and the private sector.

Seasonal to decadal

Research efforts are assessing the value of S2D forecast information for many applications, and initiatives such as the WMO’s Global Seasonal Climate Update⁵ and Annual to Decadal Climate Update (Kushnir *et al.* 2019) are making such information more widely available. However, consultation with decision makers is essential in order to tailor forecast information to the needs and expectations of users.

Fisheries management is one application for which S2D forecast information holds promise (Tommasi *et al.* 2017). This is due to reasonable skill for ocean prediction in regions of interest, coupled with strong influences of S2D climate variability on fish populations. Case studies

⁵ <https://public.wmo.int/en/our-mandate/climate/global-seasonal-climate-update>

821 employing fisheries management decision frameworks have shown that SST forecast information
822 can potentially increase fishery yields while reducing the risk of population collapse from
823 combined effects of environmental factors and overfishing. However, significant challenges
824 remain for fully realizing this potential. These include a need for improved initialization and
825 reduced model errors in key ocean regions such as the US Northeast continental shelf, dynamical
826 downscaling in cases where important spatial scales are not resolved by global models, and
827 sufficiently accurate observational data for hindcast verification on these scales. In addition,
828 incorporating biogeochemistry and marine ecosystem components into prediction systems will
829 expand their potential capabilities, while posing additional verification challenges.

830

831 Another current focus of application-oriented research is water management. Global climate
832 prediction models have been shown to have skill in predicting the next winter season's snowpack
833 throughout much of the western US, where spring snowmelt is an essential water resource
834 (Kapnick et al. 2018; Sospedra-Alfonso et al. 2016b). Because temperature influences snowmelt
835 and runoff efficiency, skill in seasonal temperature forecasts can provide improved skill for
836 seasonal water supply forecasts in this region (Lehner et al. 2017). Seasonal forecast skill has also
837 been demonstrated for monsoon rainfall (e.g., Jain et al. 2019) and drought (Hao et al. 2018) with
838 potential to inform water management decisions in many regions of the globe. Decadal forecasts
839 potentially can meet planning horizon needs but currently are less familiar to water managers
840 than seasonal forecasts and long-term climate projections. Efforts to apply decadal climate
841 information for water management decisions have included assessing the role of decadal modes
842 of variability, and using statistically downscaled decadal predictions as hydrological model inputs.

Developing information that is credible and compatible with existing decision frameworks is an important consideration (Towler et al. 2018).

Additional sectors for which S2D forecasts are being assessed for decision making include agriculture (Klemm and McPherson, 2017), energy (demand & wind power generation, Clark et al. 2017; Lledó et al. 2019), tropical cyclone (Bergman et al. 2019) and coastal flooding (Widlansky et al. 2017) preparedness, Arctic marine transportation (Stephenson and Pincus 2018), wildfire risk (Turco et al. 2019), and food security (Funk et al. 2019).

Initiatives to develop and deliver climate forecast information range in scale from international, regional and national (e.g., Marotzke et al. 2016), to individual users, all of which aim to provide forecast information having practical value for decision makers. In all cases, it is crucially important that uncertainties are adequately quantified and conveyed in order to avoid any false sense of certainty and to build trust in forecast information providers, although sometimes this requires overcoming a preference of users for deterministic information. Additional considerations are that expectations of users need to be conditioned to generally modest levels of skill, but that this information can nonetheless be advantageous when applied consistently in the long term. The likelihood that climate forecast information gets used increases when efforts are made to build relationships with potential users, and dialogs are opened to enable forecast products to be co-designed (Kolstad et al. 2019).

Cross-cutting issues in S2S and S2D prediction.

865 *Initialization shock and model error*

866 Model biases are an endemic modeling issue that is common across S2S and S2D prediction
867 time scales and influence all aspects of the prediction systems – complicating ingestion of
868 assimilated observations, degrading skill, and necessitating post-processing steps such as bias
869 correction and calibration for product development and delivery.

870

871 Model biases begin to develop on fast time scales and lead to drifts from the climate
872 represented by the initial conditions to that of a model’s biased equilibrium state. It has been
873 extremely hard to understand the mechanisms behind these drifts, and further, pathways for
874 their diagnosis are not clear although some progress is being made (Sanchez-Gomez et al. 2015;
875 Shonk et al. 2018; Voldoire et al. 2019). Such difficulties arise due to non-linear interaction
876 between various physical processes that are parameterized, and because biases could be non-
877 local in their origin. Long time scales before models’ equilibrium states are attained make
878 understanding the causes of drifts even harder. The Boulder meeting recognized that the
879 S2S/S2D prediction community needs to pay particular attention to developing pathways for
880 understanding the onset of model biases and put together mechanisms (such as summer
881 schools) to train the next generation of scientists with interest and expertise in climate
882 modeling and model diagnostics.

883

884 Initialization shocks that arise from imbalances in initial states with respect to the formulation
885 of the model and can be caused by limitations of observations and data assimilation as well as
886 model biases were also recognized as a major issue, particularly in the context of decadal

887 predictions. Initialization shocks result in the degradation of initial information that may be the
888 primary source of predictability for the subsequent forecast. Even after considerable research
889 and investment in decadal predictions it is still not clear what may be best approaches, such as
890 between full field vs. anomaly initialization, to retain predictive information in the initial state
891 while minimizing the influence of initial shocks on the subsequent forecast. The continuing
892 prominence of model drift and initial shocks as important issues reinforces a long held
893 sentiment that these are outstanding problems that need to be studied more systematically if
894 progress in translating inherent predictability into prediction skill is to be made.

896 *S2S and S2D research interactions*

897 The examples of interaction among modes of variability across S2S and S2D time scales noted
898 earlier emphasize the fact that continued interaction and communication across the S2S and
899 S2D research communities will be important to make progress. Furthering our understanding of
900 time-scale interactions will require investments in process level understanding of these
901 phenomena and will not only benefit our understanding about their lower-frequency variations
902 but will also contribute to improved process level diagnostics of model simulations. Better
903 understanding of time-scale interactions is likely to require the use of a hierarchy of models,
904 such as simple linear models to investigate the characteristics of tropical-extratropical
905 interactions (including their influence on storm tracks), to diagnose possible causes for errors in
906 their representation in complex GCMs (Dias et al. 2019).

908 Another aspect of research interactions across time scales is quantifying the fidelity of models
909 in S2S and S2D prediction as well as projections of climate on longer time scales based on their
910 simulation and prediction of shorter time-scale phenomena. The advantage of such an
911 approach is that much larger samples for predictions of shorter time-scale phenomena are
912 available, and an assessment of the reliability of such predictions can be used to build
913 confidence in prediction on longer time-scales. Theoretical basis for extrapolating the reliability
914 of forecasts across different time scales may also require the use of a hierarchy of models
915 (Weisheimer and Palmer 2014; Christensen and Berner 2019).

916

917 *Research and operations*

918 Post-processing to improve forecast quality is an important area of research that bears directly
919 on operational activities. Post-processing is necessary because biases in forecasts can be as
920 large as the predicted signal, and therefore require the use of bias correction and calibration
921 techniques to adjust real-time predictions before their delivery to the users. These
922 requirements are shared across sub-seasonal to decadal prediction time-scales, however
923 because of different levels of experience (seasonal predictions having the longest history) the
924 opportunity for cross-community interactions was recognized. Some aspects for post-
925 processing are specific to time-scale, for example, bias correction for decadal predictions
926 requires methods to account for the non-stationarity of climate, and research needs to develop
927 such methods were stressed.

928

929 Necessity for post-processing requires an extensive set of hindcasts to accompany real-time
930 predictions. Because of limited resources, decisions about hindcast period, ensemble size and
931 forecast start dates are not straightforward and call for further guidance from the research
932 community. Such questions about the operational infrastructure for long-range prediction
933 systems, including ensemble generation techniques and recommendations for harmonizing
934 hindcast and real time forecast production, provide an opportunity to link operational and
935 research communities that was highlighted during the conference.

936

937 Product development and communicating forecasts to the user community is also a common
938 thread across the S2S and S2D communities. Communication of probabilistic forecast
939 information (including confidence in the forecast based on past verifications) to users for their
940 decision making has been a challenge, and once again there is much to be gained from lessons
941 learned from the experiences of different communities. Similar challenges and opportunities
942 also exist in the context of product development that incorporate user needs based on an
943 ongoing dialog from the very start of the process. In addition, users often wish to have
944 information on smaller spatial scales than are represented in global climate models. For such
945 applications either statistical or dynamical downscaling is possible and can be effective in
946 reducing local climatological biases, although clear demonstrations that downscaling can
947 improve the skill of climate predictions remain elusive (e.g., Manzananas et al. 2018).

948

949 In summary, research needs for further development of operational infrastructure, product
950 generation and communication of probabilistic forecasts were themes often repeated during
951 the conference.

952

953 **Conclusions and the future of subseasonal to decadal prediction**

954 This paper has outlined many commonalities in the prediction of weather and climate across
955 time scales and Earth system components, and through the value cycle from basic research to
956 operational delivery.

957

958 The Earth's weather and climate is inherently chaotic and challenges the best currently
959 available modeling capabilities. There remains however untapped skill, and realizing this skill
960 will require improvements on numerous fronts. These include fundamental understanding of
961 fine-scale processes, leading toward their robust parameterization; accurately representing
962 property exchanges across Earth system components through realistic coupling limiting
963 systematic errors; sustained Earth observing systems and advanced data assimilation methods
964 enabling balanced initial conditions that avoid shocks and mitigate model drifts; and innovative
965 numerical and ensemble generation techniques to address model scalability issues. Additional
966 important avenues toward improved services include development of probabilistic information
967 for high impact weather and climate events including unprecedented extremes, and optimal
968 post-processing and data fusion to add value to multi-model ensembles, among many others.

969

These challenges are broad but so are opportunities for steady progress, ranging from curiosity-driven science to the systematic model evaluation and improvement in a collaborative and open research/operational environment.

The joint WWRP-WCRP conferences in Boulder clearly demonstrated the benefit in bringing relevant stakeholders together to cross-fertilize their experience, knowledge, respective issues and working cultures, which will surely help frame a new and vibrant research portfolio, and inspire the next generation of science leaders to ensure that society has access to the best possible weather and climate prediction science.

ACKNOWLEDGEMENTS

The International Conferences on Subseasonal to Decadal Prediction on which this paper is based were sponsored by: US CLIVAR, NSF, UCAR, NCAR and its Climate and Global Dynamics Laboratory (CGD), NOAA's Climate Variability and Predictability (CVP) and Modeling, Analysis, Predictions and Projections (MAPP) Programs, Copernicus Climate Change Service, IPSL, and WWRP/WCRP's Subseasonal-to-Seasonal (S2S) Prediction Project.

REFERENCES

- Alessandri, A., F. Catalano, M. De Felice, B. Van Den Hurk, F. Doblas Reyes, S. Boussetta, G. Balsamo, and P. A. Miller, 2017: Multi-scale enhancement of climate prediction over land by increasing the model sensitivity to vegetation variability in EC-Earth. *Climate Dyn.*, **49**, 1215–1237, <https://doi.org/10.1007/s00382-016-3372-4>.
- Al-Yaari, A., and Coauthors, 2017: Evaluating soil moisture retrievals from ESA’s SMOS and NASA’s SMAP brightness temperature datasets. *Remote Sens. Environ.*, **193**, 257–273, <https://doi.org/10.1016/j.rse.2017.03.010>.
- Ardilouze, C., and Coauthors, 2017: Multi-model assessment of the impact of soil moisture initialization on mid-latitude summer predictability. *Climate Dyn.*, **49**, 3959–3974, <https://doi.org/10.1007/s00382-017-3555-7>.
- Ardilouze, C., L. Batté, M. Déqué, E. van Meijgaard, and B. van den Hurk, 2019: Investigating the impact of soil moisture on European summer climate in ensemble numerical experiments. *Climate Dyn.*, **52**, 4011–4026, <https://doi.org/10.1007/s00382-018-4358-1>.
- Ayarzaguena, B., and Coauthors, 2018: No robust evidence of future changes in major stratospheric sudden warmings: a multi-model assessment from CCMI. *Atmos. Chem. Phys.*, **18**, 11277–11287. <http://doi.org/10.5194/acp-18-11277-2018>.
- Baggett, C., E. A. Barnes, E. Maloney, and B. Mundhenk, 2017: Advancing Atmospheric River Forecasts into Subseasonal Timescales. *Geophys. Res. Lett.*, **44**, 7528–7536, <https://doi.org/10.1002/2017GL074434>.
- Baggett, C.F., K. M. Nardi, S. J. Childs, S. N. Zito, E. A. Barnes, and E. D. Maloney, 2018: Skillful subseasonal forecasts of weekly tornado and hail activity using the Madden-Julian

1014 Oscillation. *J. Geophys. Res. Atmos.*, **123**, 661–675,
 1015 <https://doi.org/10.1029/2018JD029059>.

1016 Balmaseda, M. A., and D. Anderson, 2009: Impact of initialization strategies and observations
 1017 on seasonal forecast skill. *Geophys. Res. Lett.*, **36**, L01701,
 1018 <https://doi.org/10.1029/2008GL035561>.

1019 Balmaseda, M. A., and Coauthors, 2015: The Ocean Reanalyses Intercomparison Project (ORA-
 1020 IP). *J. Oper. Oceanogr.*, **8** (Suppl.), s80–s97,
 1021 <https://doi.org/10.1080/1755876X.2015.1022329>.

1022 Balsamo, G., and Coauthors, 2018: Satellite and in situ observations for advancing global Earth
 1023 surface modelling: A Review. *Remote Sensing*, **10**, 2038,
 1024 <https://doi.org/10.3390/rs10122038>.

1025 Barnston, A. G., M. K. Tippett, M. Ranganathan, and M. L’Heureux, 2017: Deterministic skill of
 1026 ENSO predictions from the North American Multimodel Ensemble. *Climate Dyn.*,
 1027 <https://doi.org/10.1007/s00382-017-3603-3>.

1028 Batté L., C. Ardilouze, and M. Déqué, 2018: Forecasting West African heat waves at sub-
 1029 seasonal and seasonal time scales. *Mon. Wea. Rev.*, **146**, 889–907,
 1030 <https://doi.org/10.1175/MWR-D-17-0211.1>.

1031 Bauer, P., A. Thorpe, and G. Brunet, 2015: The quiet revolution of numerical weather
 1032 prediction. *Nature*, **525**, 47–55, <https://doi.org/10.1038/nature14956>.

1033 Becker, E. J., 2017: Prediction of short-term climate extremes with a multimodel ensemble.
 1034 *Climate Extremes: Patterns and Mechanisms*, S.-Y. Wang et al. Eds., John Wiley & Sons,
 1035 Inc., 347–359, <https://doi.org/10.1002/9781119068020.ch21>.

1036 Bellucci, A., and Coauthors, 2015: Advancements in decadal climate predictability: The role of
 1037 nonoceanic drivers. *Rev. Geophys.*, **53**, 165–202,
 1038 <https://doi.org/10.1002/2014RG000473>.

1039 Bergman, D. L., L. Magnusson, J. Nilsson, and F. Vitart, 2019: Seasonal Forecasting of Tropical
 1040 Cyclone Landfall Using ECMWF’s System 4. *Wea. Forecasting*, **34**, 1239–1255,
 1041 <https://doi.org/10.1175/WAF-D-18-0032.1>.

1042 Beverley, J. D., S. J. Woolnough, L. H. Baker, S. J. Johnson, and A. Weisheimer, 2019: The
 1043 northern hemisphere circumglobal teleconnection in a seasonal forecast model and its
 1044 relationship to European summer forecast skill. *Climate Dyn.*, **52**, 3759–3771,
 1045 <https://doi.org/10.1007/s00382-018-4371-4>.

1046 Bilodeau, B., M. Carrera, A. Russell, X. Wang, and S. Belair, 2016: Impacts of SMAP data in
 1047 Environment Canada’s Regional Deterministic Prediction System. *2016 IEEE*
 1048 *International Geoscience and Remote Sensing Symposium (IGARSS)*, 5233–5236,
 1049 <https://doi.org/10.1109/IGARSS.2016.7730363>.

1050 Blockley, E. W., and K. A. Peterson, 2018: Improving Met Office seasonal forecasts of Arctic sea
 1051 ice using assimilation of CryoSat-2 thickness. *Cryosphere*, **12**, 3419–3438,
 1052 <https://doi.org/10.5194/tc-12-3419-2018>.

1053 Boer, G. J., and Coauthors, 2016: The Decadal Climate Prediction Project (DCPP) contribution to
 1054 CMIP6. *Geosci. Model Dev.*, **9**, 3751–3777, <https://doi.org/10.5194/gmd-9-3751-2016>.

1055 Booth, B. B. B., N. J. Dunstone, P. R. Halloran, T. Andrews, and N. Bellouin, 2012: Aerosols
 1056 implicated as a prime driver of twentieth-century North Atlantic climate variability.
 1057 *Nature*, **484**, 228–232, <https://doi.org/10.1038/nature10946>.

1058 Borchert, L., W. A. Müller, and J. Baehr, 2018: Atlantic Ocean Heat Transport Influences
 1059 Interannual-to-Decadal Surface Temperature Predictability in the North Atlantic Region.
 1060 *J. Clim.*, **31**, 6763–6782, <https://doi.org/10.1175/JCLI-D-17-0734.1>.

1061 Browne, P. A., P. de Rosnay, H. Zuo, A. Bennett, AND A. Dawson, 2019: Weakly Coupled Ocean–
 1062 Atmosphere Data Assimilation in the ECMWF NWP System. *Remote Sens.*, **11**, 234,
 1063 <https://doi.org/10.3390/rs11030234>.

1064 Brune, S., A. Düsterhus, H. Pohlmann, W. A. Müller, and J. Baehr, 2017: Time dependency of the
 1065 prediction skill for the North Atlantic subpolar gyre in initialized decadal hindcasts.
 1066 *Climate Dyn.*, **51**, 1947–1970, <https://doi.org/10.1007/s00382-017-3991-4>.

1067 Buckley, M. W., T. DelSole, M. S. Lozier, and L. Li, 2019: Predictability of North Atlantic sea
 1068 surface temperature and upper-ocean heat content, *J. Climate*, **32**, 3005–3023,
 1069 <https://doi.org/10.1175/JCLI-D-18-0509.1>.

1070 Butchart, N., and Coauthors, 2018: Overview of experiment design and comparison of models
 1071 participating in phase 1 of the SPARC Quasi-Biennial Oscillation initiative (QBOi). *Geosci.*
 1072 *Model Dev.*, **11**, 1009–1032, <https://doi.org/10.5194/gmd-11-1009-2018>.

1073 Butler, A. H., and Coauthors, 2016: The Climate-System Historical Forecast Project: Do
 1074 stratosphere-resolving models make better seasonal climate predictions in boreal
 1075 winter? *Quart. J. Roy. Meteor. Soc.*, **142**, 1413–1427, <https://doi.org/10.1002/qj.2743>.

1076 Butler, A., and Coauthors, 2019: Sub-seasonal Predictability and the Stratosphere. *Sub-Seasonal*
 1077 *to Seasonal Prediction*, A. W. Robertson & F. Vitart, Eds.. Elsevier, 223–241,
 1078 <https://doi.org/10.1016/B978-0-12-811714-9.00011-5>.

1079 Cai, W., and Coauthors, 2019: Pantropical climate interactions. *Science*, **363**, eaav4236,
1080 <https://doi.org/10.1126/science.aav4236>.

1081 Capotondi, A., and Coauthors, 2015: Understanding ENSO diversity. *Bull. Amer. Meteor. Soc.*,
1082 **96**, 921–938, <https://doi.org/10.1175/BAMS-D-13-00117.1>.

1083 Caron, L.-P., L. Hermonson, A. Dobbin, J. Imbers, L. Lledó L, and G. A. Vecchi, 2017: How skilful
1084 are the multi-annual forecasts of Atlantic hurricane activity? *Bull. Amer. Meteor. Soc.*,
1085 **99**, 403–413, <https://doi.org/10.1175/bams-d-17-0025.1>.

1086 Cassou, C., Y. Kushnir, E. Hawkins, A. Pirani, F. Kucharski, I. Kang, and N. Caltabiano, 2017:
1087 Decadal climate variability and predictability: Challenges and opportunities. *Bull. Amer.*
1088 *Meteor. Soc.*, **99**, 479–490, <https://doi.org/10.1175/BAMS-D-16-0286.1>.

1089 Chen, X., J. Ling, and C. Li, 2016: Evolution of the Madden–Julian oscillation in two types of El
1090 Niño. *J. Climate*, **29**, 1919–1934, <https://doi.org/10.1175/JCLI-D-15-0486.1>.

1091 Chen, Z., J. Liu, M. Song, Q. Yang, and S. Xu, 2017: Impacts of assimilating satellite sea ice
1092 concentration and thickness on Arctic sea ice prediction in the NCEP Climate Forecast
1093 System. *J. Climate*, **30**, 8429–8446, <https://doi.org/10.1175/JCLI-D-17-0093.1>.

1094 Chevallier, M., and Coauthors, 2017: Intercomparison of the Arctic sea ice cover in global
1095 ocean–sea ice reanalyses from the ORA-IP project. *Climate Dyn.*, **49**, 1107–1136,
1096 <https://doi.org/10.1007/s00382-016-2985-y>.

1097 Chevallier, M., F. Massonnet, H. Goessling, V. Guémas, and T. Jung, 2019: The Role of Sea Ice in
1098 Sub-seasonal Predictability. *Sub-Seasonal to Seasonal Prediction*, A. W. Robertson & F.
1099 Vitart, Eds., Elsevier, 201–221, <https://doi.org/10.1016/B978-0-12-811714-9.00010-3>.

1100 Chikamoto, Y., A. Timmermann, M. J. Widlansky, M. A. Balmaseda, and L. Stott, 2017: Multi-
 1101 year predictability of climate, drought, and wildfire in southwestern North America. *Sci.*
 1102 *Rep.*, **7**, 6568, <https://doi.org/10.1038/s41598-017-06869-7>.
 1103 Christensen, H. M., and J. Berner, 2019: From reliable weather forecasts to skilful climate
 1104 response: A dynamical systems approach. *Quart. J. Roy. Meteor. Soc.*, **145**, 1052-1069,
 1105 <https://doi.org/10.1002/qj.3476>.
 1106 Clark, R. T., P. E. Bett, H. E. Thornton, and A. A. Scaife, 2017: Skilful seasonal predictions for the
 1107 European energy industry. *Environ. Res. Lett.*, **12**, 12 024002,
 1108 <https://doi.org/10.1088/1748-9326/aa57ab>.
 1109 Coelho, C. A. S., M. A. F. Firpo, and F. M. de Andrade, 2018: A verification framework for South
 1110 American sub-seasonal precipitation predictions, *Meteorologische Zeitschrift*, **27**, 503 -
 1111 520, <https://doi.org/10.1127/metz/2018/0898>.
 1112 de Andrade, F. M., C. A. S. Coelho, and I. F. A. Cavalcanti, 2019: Global precipitation hindcast
 1113 quality assessment of the Subseasonal to Seasonal (S2S) prediction project models.
 1114 *Climate Dyn.*, **52**, 5451–5475, <https://doi.org/10.1007/s00382-018-4457-z>.
 1115 DeFlorio, M., D. Waliser, B. Guan, F. Ralph, and F. Vitart, 2019: Global evaluation of atmospheric
 1116 river subseasonal prediction skill. *Climate Dyn.*, **52**, 3039-3060,
 1117 <https://doi.org/10.1007/s00382-018-4309-x>
 1118 DelSole, T., and M. Tippett, 2016: Forecast comparison based on random walks. *Mon. Wea.*
 1119 *Rev.*, **144**, 615–626, <https://doi.org/10.1175/MWR-D-15-0218.1>.

1120 DelSole, T., L. Trenary, M. K. Tippett, and K. Pegion, 2017: Predictability of week-3–4 average
 1121 temperature and precipitation over the contiguous United States. *J. Climate*, **30**, 3499–
 1122 3512, <https://doi.org/10.1175/JCLI-D-16-0567.1>.
 1123 DeMott, C. A., N. P. Klingaman, and S. J. Woolnough, 2015: Atmosphere-ocean coupled
 1124 processes in the Madden-Julian oscillation. *Rev. Geophys.*, **53**, 1099–1154,
 1125 <https://doi.org/10.1002/2014RG000478>.
 1126 Dias, D. F., A. Subramanian, L. Zanna, and A. J. Miller, 2019: Remote and local influences in
 1127 forecasting Pacific SST: A linear inverse model and a multimodel ensemble study.
 1128 *Climate Dyn.*, **52**, 3183–3201, <https://doi.org/10.1007/s00382-018-4323-z>.
 1129 DiNezio and Coauthors, 2017: A 2 year forecast for a 60–80% chance of La Niña in 2017–2018.
 1130 *Geophys. Res. Lett.*, **44**, 11,624–11,635, <https://doi.org/10.1002/2017GL074904>.
 1131 DiNezio, P. N., C. Deser, Y. Okumura, and A. Karspeck, 2017: Predictability of 2-year La Niña
 1132 events in a coupled general circulation model. *Climate Dyn.*, **49**, 4237–4261,
 1133 <https://doi.org/10.1007/s00382-017-3575-3>.
 1134 Dirkson, A., W. J. Merryfield, and A. Monahan, 2017: Impacts of sea ice thickness initialization
 1135 on seasonal Arctic sea ice predictions. *J. Climate*, **30**, 1001–1017,
 1136 <https://doi.org/10.1175/JCLI-D-16-0437.1>.
 1137 Dirmeyer, P. A., and S. Halder, 2016: Sensitivity of numerical weather forecasts to initial soil
 1138 moisture variations in CFSv2. *Wea. Forecasting*, **31**, 1973–1983,
 1139 <https://doi.org/10.1175/WAF-D-16-0049.1>.

1140 Dirmeyer, P. A., and S. Halder, 2017: Application of the land–atmosphere coupling paradigm to
 1141 the operational Coupled Forecast System (CFSv2). *J. Hydrometeor.*, **18**, 85–108,
 1142 <https://doi.org/10.1175/JHM-D-16-0064.1>.

1143 Dirmeyer, P. A., S. Halder, and R. Bombardi, 2018a: On the Harvest of Predictability from Land
 1144 States in a Global Forecast Model. *J. Geophys. Res.*, **123**, 13,111–13,127,
 1145 <https://doi.org/10.1029/2018JD029103>.

1146 Dirmeyer, P. A., and Coauthors, 2018b: Verification of land–atmosphere coupling in forecast
 1147 models, reanalyses, and land surface models using flux site observations. *J.*
 1148 *Hydrometeor.*, **19**, 375–392, <https://doi.org/10.1175/JHM-D-17-0152.1>.

1149 Dirmeyer, P. A., P. Gentine, M. B. Ek, and G. Balsamo, 2019: Land Surface Processes Relevant to
 1150 S2S Prediction. *Sub-Seasonal to Seasonal Prediction*, A. W. Robertson & F. Vitart, Eds.,
 1151 Elsevier, 166–182, <https://doi.org/10.1016/B978-0-12-811714-9.00008-5>.

1152 Dobrynin, M. and Coauthors, 2018: Improved teleconnection-based dynamical seasonal
 1153 predictions of boreal winter. *Geophys. Res. Lett.*, **45**, 3605– 3614,
 1154 <https://doi.org/10.1002/2018GL077209>.

1155 Domeisen, D. I. V., C. I. Garfinkel, and A. H. Butler, 2019a: The teleconnection of El Niño
 1156 Southern Oscillation to the stratosphere. *Rev. Geophys.*,
 1157 <https://doi.org/10.1029/2018RG000596>.

1158 Domeisen, D. I. V., and Coauthors, 2019b: The role of stratosphere-troposphere coupling in sub-
 1159 seasonal to seasonal prediction. Part I: Predictability arising from stratosphere-
 1160 troposphere coupling. *J. Geophys. Res.*, <https://doi.org/10.1029/2019JD030920>.

1161 Domeisen, D. I. V., and Coauthors, 2019c: The role of stratosphere-troposphere coupling in sub-
 1162 seasonal to seasonal prediction. Part II: Predictability arising from stratosphere-
 1163 troposphere coupling. *J. Geophys. Res.*, <https://doi.org/10.1029/2019JD030923>.

1164 Dorigo, W. A., and Coauthors, 2011: The International Soil Moisture Network: a data hosting
 1165 facility for global in situ soil moisture measurements. *Hydrol. Earth Syst. Sci.*, **15**, 1675–
 1166 1698, <https://doi.org/10.5194/hess-15-1675-2011>.

1167 Dunstone, N. J., D. M. Smith, A. Scaife, L. Hermanson, R. Eade, N. Robinson, M. Andrews, and J.
 1168 Knight, 2016: Skilful predictions of the winter North Atlantic Oscillation one year ahead.
 1169 *Nat. Geosci.*, **9**, 809–814, <https://doi.org/10.1038/ngeo2824>.

1170 Dunstone, N. J., and Coauthors, 2018: Skilful seasonal predictions of summer European rainfall.
 1171 *Geophys. Res. Lett.*, **45**, 3246–3254, <https://doi.org/10.1002/2017GL076337>.

1172 Düsterhus, A., 2019: Seasonal statistical-dynamical prediction of the North Atlantic Oscillation
 1173 by probabilistic post-processing. *Nonlin. Processes Geophys. Discuss.*,
 1174 <https://doi.org/10.5194/npg-2019-50>.

1175 Eade, R., D. Smith, A. Scaife, E. Wallace, N. Dunstone, L. Hermanson, and N. Robinson, 2014: Do
 1176 seasonal-to-decadal climate predictions underestimate the predictability of the real
 1177 world? *Geophys. Res. Lett.*, **41**, 5620–5628, <https://doi.org/10.1002/2014GL061146>.

1178 Entekhabi, D., and Coauthors, 2010: The Soil Moisture Active and Passive (SMAP) Mission. *Proc.*
 1179 *IEEE*, **98**, 704–716, <https://doi.org/10.1109/JPROC.2010.2043918>.

1180 Funk, C. and Coauthors, 2019: Recognizing the Famine Early Warning Systems Network: Over 30
 1181 years of drought early warning science advances and partnerships promoting global

1182 food security. *Bull. Amer. Meteor. Soc.*, **100**, 1011–1027, <https://doi.org/10.1175/BAMS->
1183 D-17-0233.1.

1184 Ferranti, L., L. Magnusson, F. Vitart, and D. S. Richardson, 2018: How far in advance can we
1185 predict changes in large-scale flow leading to severe cold conditions over Europe?
1186 *Quart. J. Roy. Meteor. Soc.*, **144**, 1788–1802, <https://doi.org/10.1002/qj.3341>.

1187 Flato, G. M., 2011: Earth System Models: An overview. *Wiley Interdiscip. Rev.: Climate Change*, **2**,
1188 783–800, <https://doi.org/10.1002/wcc.148>.

1189 Fujii, Y., and Coauthors, 2015: Evaluation of the Tropical Pacific Observing System from the
1190 ocean data assimilation perspective. *Quart. J. Roy. Meteor. Soc.*, **141**, 2481–2496,
1191 <https://doi.org/10.1002/qj.2579>.

1192 Garfinkel, C. I., D. W. Waugh, and E. Gerber, 2013: Effect of tropospheric jet latitude on
1193 coupling between the stratospheric polar vortex and the troposphere. *J. Climate*, **26**,
1194 2077–2095, <https://doi.org/10.1175/JCLI-D-12-00301.1>.

1195 Garfinkel, C. I., and C. Schwartz, 2017: MJO-related tropical convection anomalies lead to more
1196 accurate stratospheric vortex variability in subseasonal forecast models. *Geophys. Res.*
1197 *Lett.*, **44**, 10 054–10 062, <https://doi.org/10.1002/2017GL074470>.

1198 Garfinkel, C. I., C. Schwartz, D. I. V. Domeisen, S.-W. Son, A. H. Butler, and I. P. White, 2018:
1199 Extratropical Atmospheric Predictability From the Quasi-Biennial Oscillation in
1200 Subseasonal Forecast Models. *J. Geophys. Res.*, **123**, 7855–7866,
1201 <https://doi.org/10.1029/2018JD028724>.

1202 Gleixner, S., N. S. Keenlyside, T. D. Demissie, F. Counillon, Y. Wang, and E. Viste, 2017: Seasonal
 1203 predictability of Kiremt rainfall in coupled general circulation models. *Environ. Res. Lett.*,
 1204 **12**, 114016, <https://doi.org/10.1088/1748-9326/aa8cfa>.
 1205 Graham, R. J., and Coauthors, 2011: Long-range forecasting and the Global Framework for
 1206 Climate Services. *Climate Res.*, **47**, 47–55, <https://doi.org/10.3354/cr00963>.
 1207 Hackert, E., R. M. Kovach, A. J. Busalacchi, and J. Ballabrera-Poy, 2019: Impact of Aquarius and
 1208 SMAP satellite sea surface salinity observations on coupled El Niño/Southern Oscillation
 1209 forecasts. *J. Geophys. Res.*, <https://doi.org/10.1029/2019JC015130>.
 1210 Hansen, F., R. J. Greatbatch, G. Gollan, T. Jung, and A. Weisheimer, 2017: Remote control of
 1211 North Atlantic Oscillation predictability via the stratosphere. *Quart. J. Roy. Meteor. Soc.*,
 1212 **143**, 706–719, <https://doi.org/10.1002/qj.2958>.
 1213 Hao, Z., V.P. Singh, and Y. Xia, 2018: Seasonal drought prediction: advances, challenges, and
 1214 future prospects. *Rev. Geophys.*, **56**, 108–141, <https://doi.org/10.1002/2016RG000549>.
 1215 Hazeleger, W., B. J. J. M. van den Hurk, E. Min, G. J. van Oldenborgh, A. C. Petersen, D. A.
 1216 Stainforth, E. Vasileiadou, L. A. and Smith, 2015: Tales of future weather. *Nat. Climate*
 1217 *Change*, **5**, 107–113, <https://doi.org/10.1038/nclimate2450>.
 1218 Henderson, S. A., E. D. Maloney, and S.-W. Son, 2017: Madden–Julian oscillation
 1219 teleconnections: The impact of the basic state and MJO representation in general
 1220 circulation models. *J. Climate*, **30**, 4567–4587, [https://doi.org/10.1175/JCLI-D-16-](https://doi.org/10.1175/JCLI-D-16-0789.1)
 1221 0789.1.

1222 Infanti J. M., and B. P. Kirtman, 2019: A comparison of CCSM4 high-resolution and low-
 1223 resolution predictions for south Florida and southeast United States drought. *Climate*
 1224 *Dyn.*, **52**, 6877–6892, <https://doi.org/10.1007/s00382-018-4553-0>.
 1225 Ilyina, T., and P. Friedlingstein, 2016: Biogeochemical Cycles and Climate Change. White Paper on
 1226 the World Climate Research Programme Grand Challenge on Carbon Feedbacks in the
 1227 Climate System, 10 pp, [https://www.wcrp-](https://www.wcrp-climate.org/JSC37/Documents/BGCGC_whitepaper_submission.pdf)
 1228 [climate.org/JSC37/Documents/BGCGC_whitepaper_submission.pdf](https://www.wcrp-climate.org/JSC37/Documents/BGCGC_whitepaper_submission.pdf)
 1229 Jain, S., A. A. Scaife, and A. K. Mitra, 2019: Skill of Indian summer monsoon rainfall prediction in
 1230 multiple seasonal prediction systems. *Climate Dyn.*, **52**, 5291–5301,
 1231 <https://doi.org/10.1007/s00382-018-4449-z>.
 1232 Jeong, J.-H., H. W. Linderholm, S.-H. Woo, C. Folland, B.-M. Kim, S.-J. Kim, and D. Chen, 2012:
 1233 Impacts of Snow Initialization on Subseasonal Forecasts of Surface Air Temperature for
 1234 the Cold Season. *J. Climate*, **26**, 1956–1972, <https://doi.org/10.1175/JCLI-D-12-00159.1>.
 1235 Jia, L., and Coauthors, 2015: Improved seasonal prediction of temperature and precipitation
 1236 over land in a high-resolution GFDL climate model. *J. Climate*, **28**, 2044–2062,
 1237 <https://doi.org/10.1175/JCLI-D-14-00112.1>.
 1238 Kadow, C., S. Illing, I. Kröner, U. Ulbrich, and U. Cubasch, 2017: Decadal climate predictions
 1239 improved by ocean ensemble dispersion filtering. *J. Adv. Model. Earth Syst.*, **9**, 1138–
 1240 1149, <https://doi.org/10.1002/2016MS000787>.
 1241 Kang, S. M., Y. Shin, and S.-P. Xie, 2018: Extratropical forcing and tropical rainfall distribution:
 1242 Energetics framework and ocean Ekman advection. *npj Climate Atmos. Sci.*, **1**, 20172,
 1243 <https://doi.org/10.1038/s41612-017-0004-6>

1244 Kapnick, S. B., and Coauthors, 2018: Potential for western US seasonal snowpack prediction.
 1245 *Proc. Natl. Acad. Sci. USA*, **115**, 1180–1185, <https://doi.org/10.1073/pnas.1716760115>.
 1246 Karpechko, A.Y., 2018: Predictability of sudden stratospheric warmings in the ECMWF
 1247 extended-range forecast system. *Mon. Wea. Rev.*, **146**, 1063–1075.
 1248 <https://doi.org/10.1175/MWR-D-17-0317.1>.
 1249 Kerr, Y. H., and Coauthors, 2010: The SMOS mission: New tool for monitoring key elements of
 1250 the global water cycle. *Proc. IEEE*, **98**, 666–687,
 1251 <https://doi.org/10.1109/JPROC.2010.2043032>.
 1252 Khodri, M., and Coauthors, 2017: Tropical explosive volcanic eruptions can trigger El Niño by
 1253 cooling tropical Africa. *Nat. Commun.*, **8**, 778, [https://doi.org/10.1038/s41467-017-](https://doi.org/10.1038/s41467-017-00755-6)
 1254 [00755-6](https://doi.org/10.1038/s41467-017-00755-6).
 1255 Kidston, J., A. A. Scaife, S. C. Hardiman, D. M. Mitchell, N. Butchart, M. P. Baldwin, and L. J. Gray,
 1256 2015: Stratospheric influence on tropospheric jet streams, storm tracks and surface
 1257 weather. *Nat. Geosci.*, **8**, 433–440, <https://doi.org/10.1038/ngeo2424>.
 1258 Kim, H.-M., 2017: The impact of the mean moisture bias on the key physics of MJO propagation
 1259 in the ECMWF reforecast, *J. Geophys. Res.*, **122**, 7772–7784,
 1260 <https://doi.org/10.1002/2017JD027005>.
 1261 Kim, Y.-H., and H.-Y. Chun, 2015: Momentum forcing of the quasi-biennial oscillation by
 1262 equatorial waves in recent reanalyses. *Atmos. Phys. Chem.*, **15**, 6577–6587,
 1263 <https://doi.org/10.5194/acp-15-6577-2015>.
 1264 Kim, H., F. Vitart, and D. E. Waliser, 2018: Prediction of the Madden–Julian Oscillation: A
 1265 Review. *J. Climate*, **31**, 9425–9443, <https://doi.org/10.1175/JCLI-D-18-0210.1>.

1266 Kim, W. M., S. G. Yeager, and G. Danabasoglu, 2018: Key role of internal ocean dynamics in
 1267 Atlantic multidecadal variability during the last half century. *Geophys. Res. Lett.*, **45**,
 1268 13,449–13,457, <https://doi.org/10.1029/2018GL080474>.

1269 Kirtman, B., D. Anderson, G. Brunet, I. S. Kang, A. A. Scaife, and D. M. Smith, 2013: Predictions
 1270 from weeks to decades. *Climate Science for Serving Society*, G. R. Asrar and J. W. Hurrell,
 1271 Eds., Springer, 205–235, https://doi.org/10.1007/978-94-007-6692-1_8.

1272 Kirtman, B. P., and Coauthors, 2014: The North American Multimodel Ensemble: Phase-1
 1273 seasonal-to-interannual prediction; Phase-2 toward developing intraseasonal prediction.
 1274 *Bull. Amer. Meteor. Soc.*, **95**, 585–601, <https://doi.org/10.1175/BAMS-D-12-00050.1>.

1275 Klemm, T., and R. A. McPherson, 2017: The development of seasonal climate forecasting for
 1276 agricultural producers. *Agric. For. Meteor.*, **232**, 384–399,
 1277 <https://doi.org/10.1016/j.agrformet.2016.09.005>.

1278 Kolstad, E. W., and Coauthors, 2019: Trials, errors and improvements in co-production of
 1279 climate services. *Bull. Amer. Meteor. Soc.*, **100**, 1419–1428, [https://doi.org/10.1175/BAMS-D-](https://doi.org/10.1175/BAMS-D-18-0201.1)
 1280 [18-0201.1](https://doi.org/10.1175/BAMS-D-18-0201.1).

1281 Koster, R., and Coauthors, 2004: Regions of strong coupling between soil moisture and
 1282 precipitation. *Science*, **305**, 1138–1140, <https://doi.org/10.1126/science.1100217>.

1283 Koster, R. D., Z. Guo, P. A. Dirmeyer, R. Yang, K. Mitchell, and M. J. Puma, 2009: On the nature
 1284 of soil moisture in land surface models. *J. Climate*, **22**, 4322–4335,
 1285 <https://doi.org/10.1175/2009JCLI2832.1>.

1286 Koster, R. D., and Coauthors, 2011: The second phase of the Global Land–Atmosphere Coupling
 1287 Experiment: Soil moisture contributions to subseasonal forecast skill. *J. Hydrometeor.*,
 1288 **12**, 805–822, <https://doi.org/10.1175/2011JHM1365.1>.

1289 Koster, R. D., Y. Chang, H. Wang, and S. D. Schubert, 2016: Impacts of local soil moisture
 1290 anomalies on the atmospheric circulation and on remote surface meteorological fields
 1291 during boreal summer: A comprehensive analysis over North America. *J. Climate*, **29**,
 1292 7345–7364, <https://doi.org/10.1175/JCLI-D-16-0192.1>.

1293 Kröger, J., and Coauthors, 2018: Full-field initialized decadal predictions with the MPI Earth
 1294 System Model: An initial shock in the North Atlantic. *Clim. Dyn.*, **51**, 2593–2608,
 1295 <https://doi.org/10.1007/s00382-017-4030-1>.

1296 Kushnir, Y., and Coauthors, 2019: Towards operational predictions of the near-term climate.
 1297 *Nat. Climate Change*, **9**, 94–101, <https://doi.org/10.1038/s41558-018-0359-7>.

1298 Lee, C.-Y., S. J. Camargo, F. Vitart, A. H. Sobel, and M. K. Tippett, 2018: Sub-seasonal tropical
 1299 cyclone genesis prediction and MJO in the S2S dataset. *Wea. Forecasting*, **33**, 967–988,
 1300 <https://doi.org/10.1175/WAF-D-17-0165.1>.

1301 Lee, R., S. Woolnough, A. Charlton-Perez and F. Vitart, 2019: ENSO modulation of MJO
 1302 teleconnection to the North Atlantic & Europe. *Geophys. Res. Lett.*,
 1303 <https://doi.org/10.1029/2019GL084683>.

1304 Lehner, F., A. W. Wood, D. Llewellyn, D. B. Blatchford, A. G. Goodbody, and F. Pappenberger,
 1305 2017: Mitigating the impacts of climate nonstationarity on seasonal streamflow
 1306 predictability in the U.S. Southwest. *Geophys. Res. Lett.*, **44**, 12 208–12 217,
 1307 <https://doi.org/10.1002/2017GL076043>.

1308 Leutbecher, M., and Coauthors, 2017: Stochastic representations of model uncertainties at
 1309 ECMWF: State of the art and future vision. *Quart. J. Roy. Meteor. Soc.*, **143**, 2315–2339,
 1310 <https://doi.org/10.1002/qj.3094>.
 1311 Li, B., M. Rodell, B. F. Zaitchik, R. H. Reichle, R. D. Koster, and T. M. van Dam, 2012: Assimilation
 1312 of GRACE terrestrial water storage into a land surface model: Evaluation and potential
 1313 value for drought monitoring in western and central Europe. *J. Hydrol.*, **446–447**, 103–
 1314 115, <https://doi.org/10.1016/j.jhydrol.2012.04.035>.
 1315 Li, F., Y. J. Orsolini, N. Keenlyside, M.-L. Shen, F. Counillon, and Y. Wang, 2019: Impact of snow
 1316 initialization in subseasonal-to-seasonal winter forecasts with the Norwegian Climate
 1317 Prediction Model. *J. Geophys. Res. Atmos.*, **124**, 10,033–10,048.
 1318 <https://doi.org/10.1029/2019JD030903>.
 1319 Li, H., T. Ilyina, A. Wolfgang, A. Müller, and F. Sienz, 2016: Decadal predictions of the North
 1320 Atlantic CO₂ uptake. *Nat. Commun.*, **7**, 11076, <https://doi.org/10.1038/ncomms11076>.
 1321 Li, H. M., and T. Ilyina, 2018: Current and future decadal trends in the oceanic carbon uptake
 1322 are dominated by internal variability. *Geophys. Res. Lett.*, **45**, 916–925,
 1323 <https://doi.org/10.1002/2017GL075370>.
 1324 Li, H., T. Ilyina, W. A. Müller, and P. Landschützer., 2019: Predicting the variable ocean carbon
 1325 sink. *Science Advances*, **5**, <https://doi.org/10.1126/sciadv.aav6471>.
 1326 Li, S., and A. W. Robertson, 2015: Evaluation of submonthly precipitation forecast skill from
 1327 global ensemble prediction systems. *Mon. Wea. Rev.*, **143**, 2871–2889,
 1328 <https://doi.org/10.1175/MWR-D-14-00277.1>.

1329 Li, X., G. Gollan, R. J. Greatbatch, and R. Lu, 2018: Intraseasonal variation of the East Asian
 1330 summer monsoon associated with the Madden–Julian oscillation. *Atmos. Sci. Lett.*, **19**,
 1331 e794, <https://doi.org/10.1002/asl.794>. Lim, E.-P., and H. H. Hendon, 2017: Causes and
 1332 predictability of the negative Indian Ocean Dipole and its impact on La Niña during 2016.
 1333 *Sci. Rep.*, **7**, 12619, <https://doi.org/10.1038/s41598-017-12674-z>.
 1334 Lim, Y., S. Son, and D. Kim, 2018: MJO prediction skill of the subseasonal-to-seasonal prediction
 1335 models. *J. Climate*, **31**, 4075–4094, <https://doi.org/10.1175/JCLI-D-17-0545.1>.
 1336 Lim, Y., S.-W. Son, A. G. Marshall, H. H. Hendon, and K.-H. Seo, 2019: Influence of the QBO on
 1337 MJO prediction skill in the subseasonal-to-seasonal prediction models. *Climate Dyn.*,
 1338 <https://doi.org/10.1007/s00382-019-04719-y>.
 1339 Lin, H., J. Frederiksen, D. Straus, and C. Stan, 2019: Tropical-Extratropical Interactions and
 1340 Teleconnections. *Sub-Seasonal to Seasonal Prediction*, A. W. Robertson & F. Vitart, Eds.,
 1341 Elsevier, 143-164, <https://doi.org/10.1016/B978-0-12-811714-9.00007-3>.
 1342 Liu, X, X. Wang, and A. Kumar, 2018: Multiweek prediction skill assessment of Arctic sea ice
 1343 variability in the CFSv2. *Wea. Forecasting*, **33**, 1453–1476,
 1344 <https://doi.org/10.1175/WAF-D-18-0046.1>.
 1345 Lledó, L., V. Torralba, A. Soret, J. Ramon, and F. J. Doblas-Reyes, 2019: Seasonal forecasts of
 1346 wind power generation. *Renewable Energy*, **143**, 91-100,
 1347 <https://doi.org/10.1016/j.renene.2019.04.135>.
 1348 Long, C. S., M. Fujiwara, S. Davis, D. M. Mitchell, and C. J. Wright, 2017: Climatology and
 1349 interannual variability of dynamic variables in multiple reanalyses evaluated by the

1350 SPARC Reanalysis Intercomparison Project (S-RIP). *Atmos. Chem. Phys.*, **17**, 14 593–14
 1351 629, <https://doi.org/10.5194/acp-17-14593-2017>.

1352 Lovenduski, N., S. G. Yeager, K. Lindsay, and M. C. Long, 2019: Predicting near-term variability in
 1353 ocean carbon uptake. *Earth Syst. Dyn.*, **10**, 45-57, [https://doi.org/10.5194/esd-10-45-](https://doi.org/10.5194/esd-10-45-2019)
 1354 2019.

1355 Lowe, R., M. García-Díez, J. Ballester, J. Creswick, J.-M. Robine, F. R. Herrmann, and X. Rodó,
 1356 2016: Evaluation of an early-warning system for heat wave-related mortality in Europe:
 1357 Implications for sub-seasonal to seasonal forecasting and climate services. *Int. J. Environ.*
 1358 *Res. Public Health*, **13**, 206, <https://doi.org/10.3390/ijerph13020206>.

1359 Lu, B., A. A. Scaife, N. Dunstone, D. Smith, H.-L. Ren, Y. Liu, and R. Eade, 2017: Skillful seasonal
 1360 predictions of winter precipitation over southern China. *Environ. Res. Lett.*, **12**, 074021,
 1361 <https://doi.org/10.1088/1748-9326/aa739a>.

1362 Luo, J., S. Masson, S. K. Behera, and T. Yamagata, 2008: Extended ENSO predictions using a fully
 1363 coupled ocean–atmosphere model. *J. Climate*, **21**, 84–93,
 1364 <https://doi.org/10.1175/2007JCLI1412.1>.

1365 Luo, J.-J., G. Liu, H. Hendon, O. Alves, and T. Yamagata, 2017: Inter-basin sources for two-year
 1366 predictability of the multi-year La Niña event in 2010–2012. *Sci. Rep.*, **7**, 2276,
 1367 <https://doi.org/10.1038/s41598-017-01479-9>.

1368 Maloney, E. D., and Coauthors, 2019: Process-oriented evaluation of climate and weather
 1369 forecasting Models. *Bull. Amer. Meteor. Soc.*, [https://doi.org/10.1175/BAMS-D-18-](https://doi.org/10.1175/BAMS-D-18-0042.1)
 1370 [0042.1](https://doi.org/10.1175/BAMS-D-18-0042.1).

1371 Manzananas, R., J. M. Gutiérrez, J. Fernández, E. van Meijgaard, S. Calmanti, M. E. Magariño, A. S.
 1372 Cofiño, and S. Herrera, 2018: Dynamical and statistical downscaling of seasonal
 1373 temperature forecasts in Europe: Added value for user applications. *Climate Services*, **9**,
 1374 44-56, <https://doi.org/10.1016/j.cliser.2017.06.004>.
 1375 Mariotti, A., P. M. Ruti, and M. Rixen, 2018: Progress in subseasonal to seasonal prediction
 1376 through a joint weather and climate community effort. *NPJ Climate Atmos. Sci.*, **1**, 4,
 1377 <https://doi.org/10.1038/s41612-018-0014-z>.
 1378 Mariotti, A. and Coauthors, 2020: Windows of opportunity for skillful forecasts S2S and beyond.
 1379 *Bull. Amer. Meteor. Soc.*, <https://doi.org/10.1175/BAMS-D-18-0326.1>.
 1380 Marotzke, J., and Coauthors, 2016: MiKlip—A National Research Project on Decadal Climate
 1381 Prediction. *Bull. Amer. Meteor. Soc.*, **97**, 2379–2394, [https://doi.org/10.1175/BAMS-D-](https://doi.org/10.1175/BAMS-D-15-00184.1)
 1382 15-00184.1.
 1383 Marshall, A. G., H. H. Hendon, S.-W. Son, and Y. Lim, 2017: Impact of the quasi-biennial
 1384 oscillation on predictability of the Madden–Julian oscillation. *Climate Dyn.*, **49**, 1365–
 1385 1377, <https://doi.org/10.1007/s00382-016-3392-0>.
 1386 Maycock, A. C., and P. Hitchcock, 2015: Do split and displacement sudden stratospheric
 1387 warmings have different annular mode signatures? *Geophys. Res. Lett.*, **42**, 10943–
 1388 10951. <http://doi.org/10.1002/2015GL066754>.
 1389 McKinnon, K., A. Rhines, M. Tingly, and P. Huybers, 2016: Long-lead predictions of eastern
 1390 United States hot days from Pacific sea surface temperatures. *Nat. Geosci.*, **9**, 389–394,
 1391 <https://doi.org/10.1038/ngeo2687>.

1392 Meehl, G. A., A. Hu, and H. Teng, 2016: Initialized decadal prediction for transition to positive
 1393 phase of the interdecadal Pacific oscillation. *Nat. Commun.*, **7**, 11718,
 1394 <https://doi.org/10.1038/ncomms11718>.

1395 Menegoz, M., R. Bilbao, O. Bellprat, V. Guemas, and F. J. Doblas-Reyes, 2018: Forecasting the
 1396 climate response to volcanic eruptions: prediction skill related to stratospheric aerosol
 1397 forcing. *Environ. Res. Lett.*, **13**, 064022, <https://doi.org/10.1088/1748-9326/aac4db>.

1398 Misios, S., L. J. Gray, M. F. Knudsen, C. Karoff, H. Schmidt, and J. D. Haigh, 2019: Slowdown of
 1399 the Walker circulation at solar cycle maximum, *Proc. Natl. Acad. Sci. USA*, **116**, 7186-
 1400 7191, <https://doi.org/10.1073/pnas.1815060116>.

1401 Monerie, P.-A., J. Robson, B. Dong, and N. Dunstone, 2018: A role of the Atlantic Ocean in
 1402 predicting summer surface air temperature over North East Asia? *Climate Dyn.*, **51**, 473-
 1403 491, <https://doi.org/10.1007/s00382-017-3935-z>.

1404 Morcrette, C. J., and Coauthors, 2018: Introduction to CAUSES: Description of weather and
 1405 climate models and their near-surface temperature errors in 5 day hindcasts near the
 1406 southern Great Plains *J. Geophys. Res. Atmos.*, **123**, 2655–2683,
 1407 <https://doi.org/10.1002/2017JD027199>.

1408 Mullholland, D. P., P. Laloyaux, K. Haines, and M. Balmaseda, 2015: Origin and impact of
 1409 initialization shocks in coupled atmosphere–ocean forecasts. *Mon. Wea. Rev.*, **143**,
 1410 4631–4644, <https://doi.org/10.1175/MWR-D-15-0076.1>.

1411 Müller, W. A., and Coauthors, 2018: A higher-resolution version of the Max Planck Institute
 1412 Earth System Model (MPI-ESM1.2-HR). *J. Adv. Model. Earth Syst.*, **10**, 1383–1413.
 1413 <https://doi.org/10.1029/2017MS001217>

1414 Muñoz-Sabater, J., H. Lawrence, C. Albergel, P. de Rosnay, L. Isaksen, S. Mecklenburg, Y. Kerr,
 1415 and M. Drusch, 2019: Assimilation of SMOS brightness temperatures in the ECMWF
 1416 Integrated Forecasting System. *Quart. J. Roy. Meteor. Soc.*,
 1417 <https://doi.org/10.1002/qj.3577>.
 1418 Neddermann, N. C., W. A. Müller, M. Dobrynin, A. Düsterhus, and J. Baehr, 2019: Seasonal
 1419 predictability of European summer climate re-assessed. *Clim. Dyn.*,
 1420 <https://doi.org/10.1007/s00382-019-04678-4>.
 1421 Nie, Y. and Coauthors, 2019 : Stratospheric initial conditions provide seasonal predictability of
 1422 the North Atlantic and Arctic Oscillations. *Environ. Res. Lett.*, **14**, 034006,
 1423 <https://doi.org/10.1088/1748-9326/ab0385>.
 1424 Nishimoto, E., and S. Yoden, 2017: Influence of the stratospheric quasi-biennial oscillation on
 1425 the Madden–Julian oscillation during austral summer. *J. Atmos. Sci.*, **74**, 1105–1125,
 1426 <https://doi.org/10.1175/JAS-D-16-0205.1>.
 1427 Nnamchi, H. C., J. Li, F. Kucharski, I.-S. Kang, N. S. Keenlyside, P. Chang, and R. Farneti, 2015:
 1428 Thermodynamic controls of the Atlantic Niño. *Nat. Commun.*, **6**, 8895,
 1429 <https://doi.org/10.1038/ncomms9895>.
 1430 Nowack, P., P. Braesicke, J. Haigh, N. L. Abraham, J. Pyle, and A. Voulgarakis, 2018. Using
 1431 machine learning to build temperature-based ozone parameterizations for climate
 1432 sensitivity simulations. *Environ. Res. Lett.*, **13**, 104016, [https://doi.org/10.1088/1748-](https://doi.org/10.1088/1748-9326/aae2be)
 1433 [9326/aae2be](https://doi.org/10.1088/1748-9326/aae2be).
 1434 O'Reilly, C. H., 2018: Interdecadal variability of the ENSO teleconnection to the wintertime
 1435 North Pacific. *Climate Dyn.*, **51**, 3333–3350, <https://doi.org/10.1007/s00382-018-4081-y>.

1436 O'Reilly, C. H., T. Woollings, L. Zanna, and A. Weisheimer, 2018: The impact of tropical
 1437 precipitation on summertime Euro-Atlantic circulation via a circumglobal wave train. *J.*
 1438 *Climate*, **31**, 6481–6504, <https://doi.org/10.1175/JCLI-D-17-0451.1>.
 1439 O'Reilly, C. H., A. Weisheimer, T. Woollings, L. Gray, and D. MacLeod, 2019: The importance of
 1440 stratospheric initial conditions for winter North Atlantic Oscillation predictability and
 1441 implications for the signal-to-noise paradox. *Quart. J. Roy. Meteor. Soc.*, **145**,
 1442 <https://doi.org/10.1002/qj.3413>.
 1443 Orsolini, Y. J., R. Senan, G. Balsamo, F. J. Doblas-Reyes, F. Vitart, A. Weisheimer, A. Carrasco,
 1444 and R. E. Benestad, 2013: Impact of snow initialization on sub-seasonal forecasts. *Clim.*
 1445 *Dyn.*, **41**, 1969–1982, <https://doi.org/10.1007/s00382-013-1782-0>.
 1446 Pasternack, A., J. Bhend, M. A. Liniger, H. W. Rust, W. A. Müller, and U. Ulbrich, 2018:
 1447 Parametric decadal climate forecast recalibration (DeFoReSt 1.0), *Geosci. Model Dev.*,
 1448 **11**, 351–368, <https://doi.org/10.5194/gmd-11-351-2018>.
 1449 Patricola, C. M., S. J. Camargo, P. J. Klotzbach, R. Saravanan, and P. Chang, 2018: The influence
 1450 of ENSO flavors on western North Pacific tropical cyclone activity. *J. Climate*, **31**, 5395–
 1451 5416, <https://doi.org/10.1175/JCLI-D-17-0678.1>.
 1452 Pegion, K., and Coauthors, 2019: The Subseasonal Experiment (SubX): A multi-model
 1453 subseasonal prediction experiment. *Bull. Am. Met. Soc.*, [https://doi.org/10.1175/BAMS-](https://doi.org/10.1175/BAMS-D-18-0270.1)
 1454 [D-18-0270.1](https://doi.org/10.1175/BAMS-D-18-0270.1).
 1455 Penny, S. G., and T. M. Hamill, 2017: Coupled data assimilation for integrated Earth system
 1456 analysis and prediction. *Bull. Amer. Meteor. Soc.*, **98**, ES169–ES172,
 1457 <https://doi.org/10.1175/BAMS-D-17-0036.1>.

1458 Penny, S.G., and Coauthors, 2017: Coupled Data Assimilation for Integrated Earth System
 1459 Analysis and Prediction: Goals, Challenges and Recommendations. Technical Report.
 1460 WWRP 2017-3, World Meteorological Organization (WMO).

1461 Penny, S. G., E. Bach, K. Bhargava, C.-C. Chang, C. Da, L. Sun, and T. Yoshida, 2019: Strongly
 1462 coupled data assimilation in multiscale media: Experiments using a quasi-geostrophic
 1463 coupled model. *J. Adv. Model. Earth Syst.*, **11**, 1803–1829,
 1464 <https://doi.org/10.1029/2019MS001652>.

1465 Polkova, I., and Coauthors, 2019: Initialization and ensemble generation for decadal climate
 1466 predictions: A comparison of different methods. *J. Adv. Model. Earth Syst.*, **11**, 149–172.
 1467 <https://doi.org/10.1029/2018MS001439>.

1468 Prodhomme, C., F. Doblas-Reyes, O. Bellprat, and E. Dutra, 2016a: Impact of land-surface
 1469 initialization on sub-seasonal to seasonal forecasts over Europe. *Climate Dyn.*, **47**, 919–
 1470 935, <https://doi.org/10.1007/s00382-015-2879-4>.

1471 Prodhomme, C., L. Batté, F. Massonnet, P. Davini, O. Bellprat, V. Guemas, and F. Doblas-Reyes,
 1472 2016b: Benefits of increasing the model resolution for the seasonal forecast quality in
 1473 EC-Earth. *J. Climate*, **29**, 9141–9162, <https://doi.org/10.1175/JCLI-D-16-0117.1>.

1474 Robertson, A. W., S. J. Camargo, A. Sobel, F. Vitart, and S. Wang, 2018: Summary of workshop
 1475 on sub-seasonal to seasonal predictability of extreme weather and climate. *NPJ Climate*
 1476 *Atmos. Sci.*, **1**, 20178, <https://doi.org/10.1038/s41612-017-0009-1>.

1477 Robson, J., P. Ortega, and R. Sutton, 2016: A reversal of climatic trends in the North Atlantic
 1478 since 2005. *Nat. Geosci.*, **9**, 513–517, <https://doi.org/10.1038/ngeo2727>.

1479 Ruprich-Robert, Y., R. Msadek, F. Castruccio, S. Yeager, T. Delworth, and G. Danabasoglu, 2017:
 1480 Assessing the climate impacts of the observed Atlantic multidecadal variability using the
 1481 GFDL CM2.1 and NCAR CESM1 global coupled models. *J. Climate*, **30**, 2785–2810,
 1482 <https://doi.org/10.1175/JCLI-D-16-0127.1>.
 1483 Ruprich-Robert, Y., T. Delworth, R. Msadek, F. Castruccio, S. Yeager, and G. Danabasoglu, 2018:
 1484 Impacts of the Atlantic multidecadal variability on North American summer climate and
 1485 heat waves. *J. Climate*, **31**, 3679–3700, <https://doi.org/10.1175/JCLI-D-17-0270.1>.
 1486 Sanchez-Gomez, E., C. Cassou, Y. Ruprich-Robert, E. Fernandez, and L. Terray, 2016: Drift
 1487 dynamics in a coupled model initialized for decadal forecasts. *Climate Dyn.*, **46**, 1819–
 1488 1840, <https://doi.org/10.1007/s00382-015-2678-y>.
 1489 Santanello, J. A., and Coauthors, 2018: Land-Atmosphere Interactions: The LoCo Perspective.
 1490 *Bull. Amer. Meteor. Soc.*, **99**, 1253–1272, <https://doi.org/10.1175/BAMS-D-17-0001.1>.
 1491 Saravanan, R., and P. Chang, 2019: Midlatitude mesoscale ocean-atmosphere interaction and its
 1492 relevance to S2S prediction. *Sub-Seasonal to Seasonal Prediction*, A. W. Robertson & F.
 1493 Vitart, Eds., Elsevier, 183– 200, <https://doi.org/10.1016/B978-0-12-811714-9.00009-7>.
 1494 Scaife, A. A., and Coauthors, 2014a: Predictability of the quasi-biennial oscillation and its
 1495 northern winter teleconnection on seasonal to decadal timescales. *Geophys. Res. Lett.*,
 1496 **41**, 1752–1758, <https://doi.org/10.1002/2013GL059160>.
 1497 Scaife, A., and Coauthors, 2014b: Skillful long-range predictions of European and North
 1498 American winters. *Geophys. Res. Lett.*, **41**, 2514–2519,
 1499 <https://doi.org/10.1002/2014GL059637>.

1500 Scaife, A. A., and Coauthors, 2017: Tropical rainfall, Rossby waves and regional winter climate
 1501 predictions. *Quart. J. Roy. Meteor. Soc.*, **143**, 1–11, <https://doi.org/10.1002/qj.2910>.

1502 Scaife, A. A., and D. Smith, 2018: A signal to noise paradox in climate science. *npj Clim. Atmos.*
 1503 *Sci.*, **1**, 28, <https://doi.org/10.1038/s41612-018-0038-4>.

1504 Scaife, A. A., and Coauthors, 2019: Does increased atmospheric resolution improve seasonal
 1505 climate predictions? *Atmos. Sci. Lett.*, **20**, e922, <https://doi.org/10.1002/asl.922>.

1506 Schuster, M., and Coauthors, 2019: Improvement in the decadal prediction skill of the northern
 1507 hemisphere extra-tropical winter circulation through increased model resolution, *Earth*
 1508 *Syst. Dynam. Discuss.*, <https://doi.org/10.5194/esd-2019-18>.

1509 Sheen, K. L., D. M. Smith, N. J. Dunstone, R. Eade, D. P. Rowell, and M. Vellinga, 2017: Skilful
 1510 prediction of Sahel summer rainfall on inter-annual and multi-year timescales. *Nat.*
 1511 *Commun.*, **8**, 14966, <https://doi.org/10.1038/ncomms14966>.

1512 Shen, M.-L., N. Keenlyside, F. Selten, W. Wiegner, and G. S. Duane, 2016: Dynamically
 1513 combining climate models to “supermodel” the tropical Pacific. *Geophys. Res. Lett.*, **43**,
 1514 359–366, <https://doi.org/10.1002/2015GL066562>.

1515 Shonk, J. K., E. Guilyardi, T. Toniazzo, S. J. Woolnough, and T. Stockdale, 2018: Identifying causes
 1516 of western Pacific ITCZ drift in ECMWF System 4 hindcasts. *Climate Dyn.*, **50**, 939–954,
 1517 <https://doi.org/10.1007/s00382-017-3650-9>.

1518 Shonk, J. K. P., T. D. Demissie, and T. Toniazzo, 2019: A double ITCZ phenomenology of wind
 1519 errors in the equatorial Atlantic in seasonal forecasts with ECMWF models. . *Atmos.*
 1520 *Chem. Phys.*, **19**, 11383–11399, <https://doi.org/10.5194/acp-19-11383-2019>.

1521 Sigmond, M., J. F. Scinocca, V. V. Kharin, and T. G. Shepherd, 2013: Enhanced seasonal forecast
 1522 skill following stratospheric sudden warmings. *Nat. Geosci.*, **6**, 98–102,
 1523 doi:<https://doi.org/10.1038/ngeo1698>.

1524 Simpson, I. R., P. Hitchcock, R. Seager, Y. Wu, and P. Callaghan, 2018: The downward influence
 1525 of uncertainty in the Northern Hemisphere stratospheric polar vortex response to
 1526 climate change. *J. Climate*, **31**, 6371– 6391. <https://doi.org/10.1175/JCLI-D-18-0041.1>.

1527 Smith, D. M., and Coauthors, 2016: Role of the volcanic and anthropogenic aerosols in the
 1528 recent global surface warming slowdown. *Nat. Climate Change*, **6**, 936–940,
 1529 <https://doi.org/10.1038/nclimate3058>.

1530 Smith, D. M., and Coauthors, 2019: Robust skill of decadal climate predictions, *npj Clim. Atmos.*
 1531 *Sci.*, **2**, 13, <https://doi.org/10.1038/s41612-019-0071-y>.

1532 Sospedra-Alfonso, R., L. Mudryk, W. J. Merryfield and C. Derksen, 2016a: Representation of
 1533 snow in the Canadian Seasonal to Interannual Prediction System: Part I. Initialization. *J.*
 1534 *Hydrometeorology*, **17**, 1467–1488, <https://doi.org/10.1175/JHM-D-14-0223.1>.

1535 Sospedra-Alfonso, R., W. J. Merryfield and V. V. Kharin, 2016b: Representation of snow in the
 1536 Canadian Seasonal to Interannual Prediction System: Part II. Potential predictability and
 1537 hindcast skill. *J. Hydrometeorology*, **17**, 2511–2535, [https://doi.org/10.1175/JHM-D-16-](https://doi.org/10.1175/JHM-D-16-0027.1)
 1538 [0027.1](https://doi.org/10.1175/JHM-D-16-0027.1).

1539 Stan, C., and D. M. Straus, 2019: The impact of cloud representation on the sub-seasonal
 1540 forecasts of atmospheric teleconnections and preferred circulation regimes in the
 1541 northern hemisphere. *Atmosphere-Ocean*, **57**, 233–248,
 1542 <https://doi.org/10.1080/07055900.2019.1590178>.

1543 Stephenson, S. R., and R. Pincus, 2018: Challenges of sea-ice prediction for Arctic marine policy
 1544 and planning. *J. Borderlands Studies*, **33**, 255–272,
 1545 <https://doi.org/10.1080/08865655.2017.1294494>.
 1546 Strommen K., and T. N. Palmer, 2019: Signal and noise in regime systems: A hypothesis on the
 1547 predictability of the North Atlantic Oscillation. *Quart. J. Roy. Meteor. Soc.*, **145**, 147–163,
 1548 <https://doi.org/10.1002/qj.3414>.
 1549 Stone, K. A., S. Solomon, D. E. Kinnison, C. F. Baggett, and Elizabeth A. Barnes, 2019: Prediction
 1550 of Northern Hemisphere regional surface temperatures using stratospheric ozone
 1551 information. *J. Geophys. Res.*, **124**, 5922–5933, <https://doi.org/10.1029/2018JD029626>.
 1552 Swingedouw, D., J. Mignot, P. Ortega, M. Khodri, M. Menegoz, C. Cassou, and V. Hanquiez,
 1553 2017: Impact of explosive volcanic eruptions on the main climate variability modes.
 1554 *Global Planet. Change*, **150**, 24–45, <https://doi.org/10.1016/j.gloplacha.2017.01.006>.
 1555 Strazzo, S., D. C. Collins, A. Schepen, Q. J. Wang, E. Becker, and L. Jia, 2019: Application of a
 1556 hybrid statistical-dynamical system to seasonal prediction of North American
 1557 temperature and precipitation. *Mon. Wea. Rev.*, **147**, 607–625,
 1558 <https://doi.org/10.1175/MWR-D-18-0156.1>.
 1559 Taguchi, M., 2018: Comparison of Subseasonal-to-Seasonal Model Forecasts for Major
 1560 Stratospheric Sudden Warmings. *J. Geophys. Res.*, **123**, 10,231–10,247,
 1561 <https://doi.org/10.1029/2018JD028755>.
 1562 Takahashi, C., and M. Watanabe, 2016: Pacific trade winds accelerated by aerosol forcing over
 1563 the past two decades. *Nat. Climate Change*, **6**, 768–772,
 1564 <https://doi.org/10.1038/nclimate2996>.

1565 Takaya, Y., 2019: Forecast System Design, Configuration, and Complexity. *Sub-Seasonal to*
 1566 *Seasonal Prediction*, A. W. Robertson & F. Vitart, Eds., Elsevier, 93-117,
 1567 <https://doi.org/10.1016/B978-0-12-811714-9.00012-7>.
 1568 Teng, H., G. Branstator, A. B. Tawfik, and P. Callaghan, 2019: Circumglobal response to
 1569 prescribed soil moisture over North America. *J. Climate*, **32**, 4525– 4545.
 1570 <https://doi.org/10.1175/JCLI-D-18-0823.1>
 1571 Tommasi, D. and Coauthors, 2017: Managing living marine resources in a dynamic environment:
 1572 The role of seasonal to decadal climate forecasts. *Prog. Oceanogr.*, **152**, 15–49,
 1573 <https://doi.org/10.1016/j.pocean.2016.12.011>.
 1574 Toniazzo, T., and S. Koseki, 2018: A methodology for anomaly coupling in climate simulation. *J.*
 1575 *Adv. Model. Earth Syst.*, **10**, 2061–2079, <https://doi.org/10.1029/2018MS001288>.
 1576 Toure, A. M., R. H. Reichle, B. A., Forman, A., Getirana, and G. J. M. De Lannoy, 2018.
 1577 Assimilation of MODIS snow cover fraction observations into the NASA Catchment land
 1578 surface model. *Remote Sensing*, **10**, 316. <https://doi.org/10.3390/rs10020316>.
 1579 Towler, E., D. PaiMazumder, and J. Done, 2018: Toward the application of decadal climate
 1580 predictions. *J. App. Meteorol. Climatol.*, **57**, 555-568, [https://doi.org/10.1175/JAMC-D-](https://doi.org/10.1175/JAMC-D-17-0113.1)
 1581 [17-0113.1](https://doi.org/10.1175/JAMC-D-17-0113.1).
 1582 Tripathi, O. P., and Coauthors, 2015: The predictability of the extratropical stratosphere on
 1583 monthly time-scales and its impact on the skill of tropospheric forecasts. *Quart. J. Roy.*
 1584 *Meteor. Soc.*, **141**, 987–1003, <https://doi.org/10.1002/qj.2432>.
 1585 Turco, M., R. Marcos-Matamorosa, X. Castro, E. Canyameras, and M. C. Llasat, 2019: Seasonal
 1586 prediction of climate-driven fire risk for decision-making and operational applications in

1587 a Mediterranean region. *Sci. Total Environ.*, **676**, 577–583,
 1588 <https://doi.org/10.1016/j.scitotenv.2019.04.296>.
 1589 Uotila, P., and Coauthors, 2019: An assessment of ten ocean reanalyses in the polear regions.
 1590 *Climate Dyn.*, **52**, 1613–1650, <https://doi.org/10.1007/s00382-018-4242-z>.
 1591 Vigaud, N., A. Robertson, and M. Tippett, 2017: Multimodel ensembling of subseasonal
 1592 precipitation forecasts over North America. *Mon. Wea. Rev.*, **145**, 3913–3928,
 1593 <https://doi.org/10.1175/MWR-D-17-0092.1>.
 1594 Vitart, F., 2017: Madden–Julian oscillation prediction and teleconnections in the S2S database.
 1595 *Quart. J. Roy. Meteor. Soc.*, **143**, 2210–2220, <https://doi.org/10.1002/qj.3079>.
 1596 Vitart, F., and M. Balmaseda, 2017: Impact of sea surface temperature biases on extended-
 1597 range forecasts. ECMWF Technical Memorandum 830.
 1598 Vitart, F., and Coauthors, 2017: The Subseasonal to Seasonal (S2S) Prediction project database.
 1599 *Bull. Amer. Meteor. Soc.*, **98**, 163–173, <https://doi.org/10.1175/BAMS-D-16-0017.1>.
 1600 Vitart, F., and A. W. Robertson, 2018: The sub-seasonal to seasonal prediction project (S2S) and
 1601 the prediction of extreme events. *npj Climate Atmos. Sci.*, **1**, 3,
 1602 <https://doi.org/10.1038/s41612-018-0013-0>.
 1603 Vitart, F., and A. Robertson, 2019: Introduction: Why Sub-seasonal to Seasonal Prediction (S2S)?
 1604 *Sub-Seasonal to Seasonal Prediction*, A. W. Robertson & F. Vitart, Eds., Elsevier, 3–15,
 1605 <https://doi.org/10.1016/B978-0-12-811714-9.00001-2>.
 1606 Voltaire, A., and Coauthors, 2019 : Role of wind stress in driving SST biases in the Tropical
 1607 Atlantic. *Climate Dyn.*, **53**, 3481–3504, <https://doi.org/10.1007/s00382-019-04717-0>.

1608 Volpi, D., V. Guemas, and F. J. Doblas-Reyes, 2017: Comparison of full field and anomaly
 1609 initialisation for decadal climate prediction: Towards an optimal consistency between
 1610 the ocean and sea-ice anomaly initialisation state. *Climate Dyn.*, **49**, 1181–1195,
 1611 <https://doi.org/10.1007/s00382-016-3373-3>.

1612 Wang, H. L., S. D. Schubert, R. D. Koster, and Y. Chang, 2019: Phase locking of the boreal
 1613 summer atmospheric response to dry land surface anomalies in the Northern
 1614 Hemisphere. *J. Climate*, **32**, 1081–1099, <https://doi.org/10.1175/JCLI-D-18-0240.1>.

1615 Wang, T., D. Guo, Y. Gao, H. Wang, F. Zheng, Y. Zhu, J. Miao, and Y. Hu, 2018: Modulation of
 1616 ENSO evolution by strong tropical volcanic eruptions. *Climate Dyn.*, **51**, 2433–2453,
 1617 <https://doi.org/10.1007/s00382-017-4021-2>.

1618 Wei, J., and P. A. Dirmeyer, 2019: Sensitivity of land precipitation to surface evapotranspiration:
 1619 a nonlocal perspective based on water vapor transport. *Geophys. Res. Lett.*,
 1620 <https://doi.org/10.1029/2019GL085613>.

1621 Weisheimer, A., and T. Palmer, 2014: On the reliability of seasonal climate forecasts. *J. Roy. Soc.*
 1622 *Interface*, **11**, 20131162, <https://doi.org/10.1098/rsif.2013.1162>.

1623 Weisheimer, A., D. Decremmer, D. MacLeod, C. O'Reilly, T. N. Stockdale, S. Johnson, and T. N.
 1624 Palmer, 2019: How confident are predictability estimates of the winter North Atlantic
 1625 Oscillation? *Quart. J. Roy. Meteor. Soc.*, **145**, 1–20, <https://doi.org/10.1002/qj.3446>.

1626 Weiss, M., and Coauthors, 2014: Contribution of dynamic vegetation phenology to decadal
 1627 climate predictability. *J. Climate*, **27**, 8563–8577, [https://doi.org/10.1175/JCLI-D-13-](https://doi.org/10.1175/JCLI-D-13-00684.1)
 1628 [00684.1](https://doi.org/10.1175/JCLI-D-13-00684.1).

1629 White, C. J., and Coauthors, 2017: Potential applications of subseasonal-to-seasonal (S2S)
 1630 predictions. *Meteor. Appl.*, **24**, 315–325, <https://doi.org/10.1002/met.1654>.

1631 Widlansky, M., and Coauthors, 2017: Multi-model ensemble sea level forecasts for tropical
 1632 Pacific islands. *J. Appl. Meteor. Climatol.*, **56**, 849– 862. <https://doi.org/10.1175/JAMC->
 1633 D-16-0284.1.

1634 Williams, I. N., Y. Lu, L. M. Kueppers, W. J. Riley, S. Biraud, J. E. Bagley, and M. S. Torn, 2016:
 1635 Land–atmosphere coupling and climate prediction over the US Southern Great Plains. *J.*
 1636 *Geophys. Res. Atmos.*, **121**, 12 125–12 144, <https://doi.org/10.1002/2016JD025223>.

1637 Woolnough, S. J., 2019: The Madden-Julian Oscillation. *Sub-Seasonal to Seasonal Prediction*, A.
 1638 W. Robertson & F. Vitart, Eds., Elsevier, 93-117, <https://doi.org/10.1016/B978-0-12->
 1639 811714-9.00005-X.

1640 World Meteorological Organization, 2018: Manual on the Global Data-processing and
 1641 Forecasting System: Annex IV to the WMO Technical Regulations (2018 update). WMO-
 1642 485, 119 pp., https://library.wmo.int/doc_num.php?explnum_id=5839.

1643 Xue, Y. K., and Coauthors, 2018: Spring land surface and subsurface temperature anomalies and
 1644 subsequent downstream late spring-summer droughts/floods in North America and East
 1645 Asia. *J. Geophys. Res. Atmos.*, **129**, 5001–5019, <https://doi.org/10.1029/2017JD028246>.

1646 Yang, S., Z. Li, J.-Y. Yu, X. Hu, W. Dong, and S. He, 2018: El Niño–Southern oscillation and its
 1647 impact in the changing climate. *National Science Review*, **5**, 840–857.
 1648 <https://doi.org/10.1093/nsr/nwy046>.

1649 Yang, X., and T. DelSole, 2012: Systematic comparison of ENSO teleconnection patterns
 1650 between models and observations. *J. Climate*, **25**, 425–446,
 1651 <https://doi.org/10.1175/JCLI-D-11-00175.1>.

1652 Yang, Z., J. Zhang, and L. Wu, 2019: Spring soil temperature as a predictor of summer
 1653 heatwaves over northwestern China. *Atmos. Sci. Lett.*, **20**, e887,
 1654 <https://doi.org/10.1002/asl.887>.

1655 Yeager, S., and J. I. Robson, 2017: Recent progress in understanding and predicting Atlantic
 1656 decadal climate variability. *Curr. Climate Change Rep.*, **3**, 112–127,
 1657 <https://doi.org/10.1007/s40641-017-0064-z>.

1658 Yeager, S. G., and Coauthors, 2018: Predicting near-term changes in the Earth System: A large
 1659 ensemble of initialized decadal prediction simulations using the Community Earth System
 1660 Model. *Bull. Amer. Meteor. Soc.*, **99**, 1867–1886, [https://doi.org/10.1175/BAMS-D-17-](https://doi.org/10.1175/BAMS-D-17-0098.1)
 1661 [0098.1](https://doi.org/10.1175/BAMS-D-17-0098.1).

1662 Yeh, S.-W., and Coauthors, 2018: ENSO atmospheric teleconnections and their response to
 1663 greenhouse gas forcing. *Rev. Geophys.*, **56**, 185–206,
 1664 <https://doi.org/10.1002/2017RG000568>.

1665 Yoo, C., S. Park, D. Kim, J.-H. Yoon, and H.-M. Kim, 2015: Boreal winter MJO teleconnection in the
 1666 Community Atmosphere Model version 5 with the unified convection parameterization.
 1667 *J. Climate*, **28**, 8135–8150, <https://doi.org/10.1175/JCLI-D-15-0022.1>.

1668 Yoo, C., and S.-W. Son, 2016: Modulation of the boreal wintertime Madden-Julian oscillation by
 1669 the stratospheric quasi-biennial oscillation. *Geophys. Res. Lett.*, **43**, 1392–1398,
 1670 <https://doi.org/10.1002/2016GL067762>.

1671 Yuan, X., M. R. Kaplan, and M. A. Cane, 2018: The Interconnected Global Climate System—A
 1672 Review of Tropical–Polar Teleconnections. *J. Climate*, **31**, 5765–5792,
 1673 <https://doi.org/10.1175/JCLI-D-16-0637.1>.
 1674 Yun, K. S., and A. Timmermann, 2018: Decadal monsoon–ENSO relationships reexamined.
 1675 *Geophys. Res. Lett.*, **45**, 2014–2021, <https://doi.org/10.1002/2017GL076912>.
 1676 Zampieri, L., H. F. Goessling and T. Jung, 2018: Bright prospects for Arctic sea ice prediction on
 1677 subseasonal time scales. *Geophys. Res. Lett.*, **45**, 9731– 9738,
 1678 <https://doi.org/10.1029/2018GL079394>.
 1679 Zhang, W. and B. Kirtman, 2019: Understanding the signal-to-noise paradox with a simple
 1680 Markov model. *Geophys. Res. Lett.*, <https://doi.org/10.1029/2019GL085159>.
 1681 Zhang, L., T. L. Delworth, W. Cooke, and X. Yang, 2019: Natural variability of Southern Ocean
 1682 convection as a driver of observed climate trends. *Nat. Climate Change*, **9**, 59–65,
 1683 <https://doi.org/10.1038/s41558-018-0350-3>.
 1684 Zhang, Y., T. Zou, and Y. Xue, 2019: An Arctic-Tibetan Connection on Subseasonal to Seasonal
 1685 Time Scale. *Geophys. Res. Lett.*, **46**, 2790–2799, <https://doi.org/10.1029/2018GL081476>.
 1686 Zhao, C., H.-L. Ren, R. Eade, Y. Wu, J. Wu, and C. MacLachlan, 2019: MJO modulation and its
 1687 ability to predict boreal summer tropical cyclone genesis over the northwest Pacific in
 1688 Met Office Hadley Centre and Beijing Climate Center seasonal prediction systems.
 1689 *Quart. J. Roy. Meteor. Soc.*, **145**, 1089– 1101. <https://doi.org/10.1002/qj.3478>.
 1690 Zhu, J., and Coauthors, 2015: ENSO prediction in Project Minerva: Sensitivity to atmospheric
 1691 horizontal resolution and ensemble size. *J. Climate*, **28**, 2080–2095,
 1692 <https://doi.org/10.1175/JCLI-D-14-00302.1>.

SIDEBAR 1:

Hindcast and forecast quality assessment (or, “the unexamined life is not worth living”).

Besides helping to inform decision making, the careful assessment of forecast quality is critical to guiding forecasting improvements, but has many and varied considerations. Simply answering the question “is this forecast better than that one?” is complicated, as the appropriate skill metric or means for comparison is not always obvious. Some recent studies have focused on newer statistical methods for comparing one forecast to another. One relatively simple approach is the random walk test (DelSole and Tippett 2016), illustrated in Fig. SB1. This method is applicable to a wide range of measures such as a score based on the earth mover’s distance metric (Düsterhus 2019), while also being robust and fair in its discrimination.

The utility of forecast assessment can be illustrated through two very different applications of seasonal forecasts: sea-ice and hurricanes. The assessment of seasonal sea ice forecasts is complicated by the low quality of sea-ice observations, but nevertheless reveals that initializing sea-ice thickness using observational data sets generally improves the prediction of Arctic sea-ice extent and edges (Blockley et al. 2018). Comparison of multi-annual forecasts of Atlantic hurricane activity obtained by direct tracking of storms in decadal hindcasts and through a hybrid approach combining predicted SSTs and observed statistical relations finds that each approach is skillful, especially hybrid forecasts based on a SST index for AMV (Caron et al. 2018).

A robust assessment of model performance should include the model's simulation of climate modes and teleconnection patterns such as ENSO, MJO and NAO, since they are major sources of predictability and errors representing their patterns or strength (e.g., Yang and DelSole 2012; Vitart 2017) can degrade skill in affected regions (Gleixner et al. 2017; Lu et al. 2017). In cases where modeled teleconnection patterns are imperfect, forecast skill may be improved by means of statistical methods that use model forecasts of relevant climate modes such as ENSO as predictors (e.g., Strazzo et al. 2019). It remains desirable, however, for models to improve so that their simulated teleconnection patterns are sufficiently realistic that such corrections are not needed.

SIDEBAR 2:

Frontiers in Earth system prediction.

New frontiers in S2D prediction have been enabled by Earth system models (ESMs, Flato 2011) that represent the carbon and other biogeochemical cycles in addition to the physical climate system. These frontiers include prediction of ocean and land carbon sinks and biogeochemistry and their important contribution to global carbon storage, as well as ecosystem services. The world's oceans are a fundamental regulator of global carbon storage and variability. The strength of ocean carbon uptake, together with uptake of carbon by the land, determines the fraction of anthropogenic emissions remaining in the atmosphere, and hence modulates present and future warming. Observed global mean ocean carbon uptake shows variability on decadal time scales that can be represented by ESMs in which physical climate variables are assimilated (H. Li et al. 2019).

1736

1737 ESM simulations indicate that internal variability of the ocean carbon uptake on decadal
1738 timescales is as large as the forced climate change trend (Li and Ilyina 2018), pointing to the
1739 potential importance and utility of decadal carbon cycle predictions. Decadal predictions from a
1740 number of ESMs are assessing the predictability of the ocean and land carbon sinks and other
1741 ocean tracers such as dissolved oxygen. These forecasts are part of the Decadal Climate
1742 Prediction Project (Boer et al. 2016) and international programs such as the World Climate
1743 Research Programme's Grand Challenge on Carbon Feedbacks (Ilyina and Friedlingstein 2016).
1744 Initial results based on individual models have demonstrated potential predictive skill, assessed
1745 through verification against the assimilating reconstructions that provide initial conditions, for
1746 ocean carbon uptake in certain regions such as the North Atlantic reaching up to 7 or more
1747 years (Li et al. 2016; Lovenduski et al. 2019), and skill in predicting actual variations estimated
1748 from observations (Fig. SB2) has been demonstrated (Li et al. 2019).

1749 ESM-based studies also point to the drivers of this predictability. Air-sea CO₂ flux mainly varies
1750 due to pCO₂ changes in the ocean. While thermal influences on pCO₂ play a role in shorter term
1751 predictability, predictability beyond 3 years is maintained mainly by nonthermal influences of
1752 ocean circulation and biological modification of surface dissolved inorganic carbon and
1753 alkalinity (Li et al. 2019; Lovenduski et al. 2019).

1754

1755 Investigations in progress are finding potential for multi-annual prediction of additional
1756 biogeochemical fields such as net primary productivity and interior dissolved oxygen
1757 concentrations. In addition, potential predictability and skill for terrestrial carbon uptake have

1758 also begun to be assessed, with promising initial results (N. Lovenduski 2019, personal
1759 communication). These examples demonstrate that skillful multi-year prediction is likely
1760 achievable for biogeochemical and ecological Earth system components, and open prospects
1761 for the utilization of such information although significant challenges including the paucity of
1762 long term observational data for initialization and verification will need to be overcome.

1763

1764

1765

1766

1767

1768

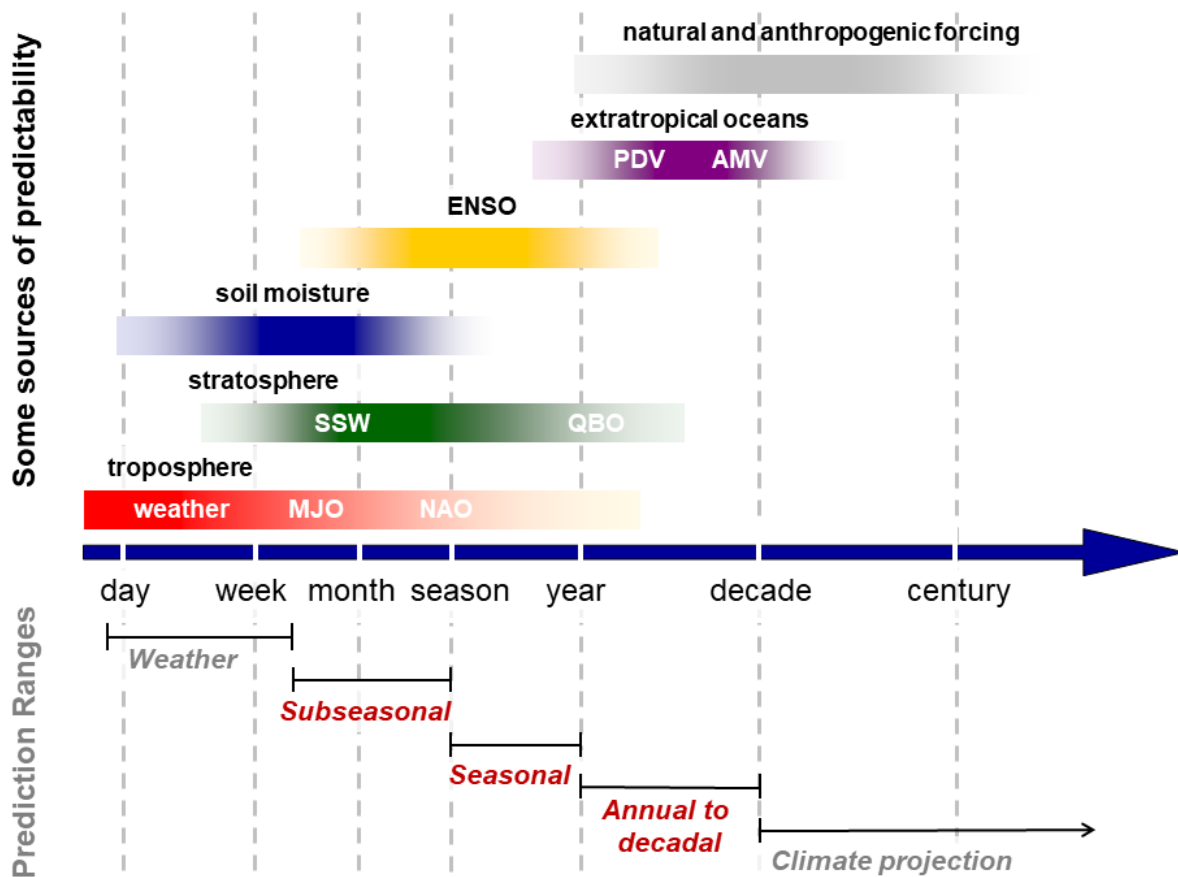


Fig. 1. Schematic depiction of temporal ranges and sources of predictability for weather and climate prediction. The subseasonal range encompasses the S2S time scales, and the seasonal and annual-to-decadal ranges the S2D time scales, that are considered in this paper. Longer multi-decadal and centennial ranges derive predictability mainly from forcing scenarios rather than initial conditions, and are typically represented through climate projections originating from historical simulations begun in preindustrial times rather than predictions initialized from more recent observation-based climate states. Some important sources of predictability and approximate time scales over which they are most influential on surface climate are indicated in the upper portion of the figure; acronyms are defined and associated phenomena are discussed in the main text.

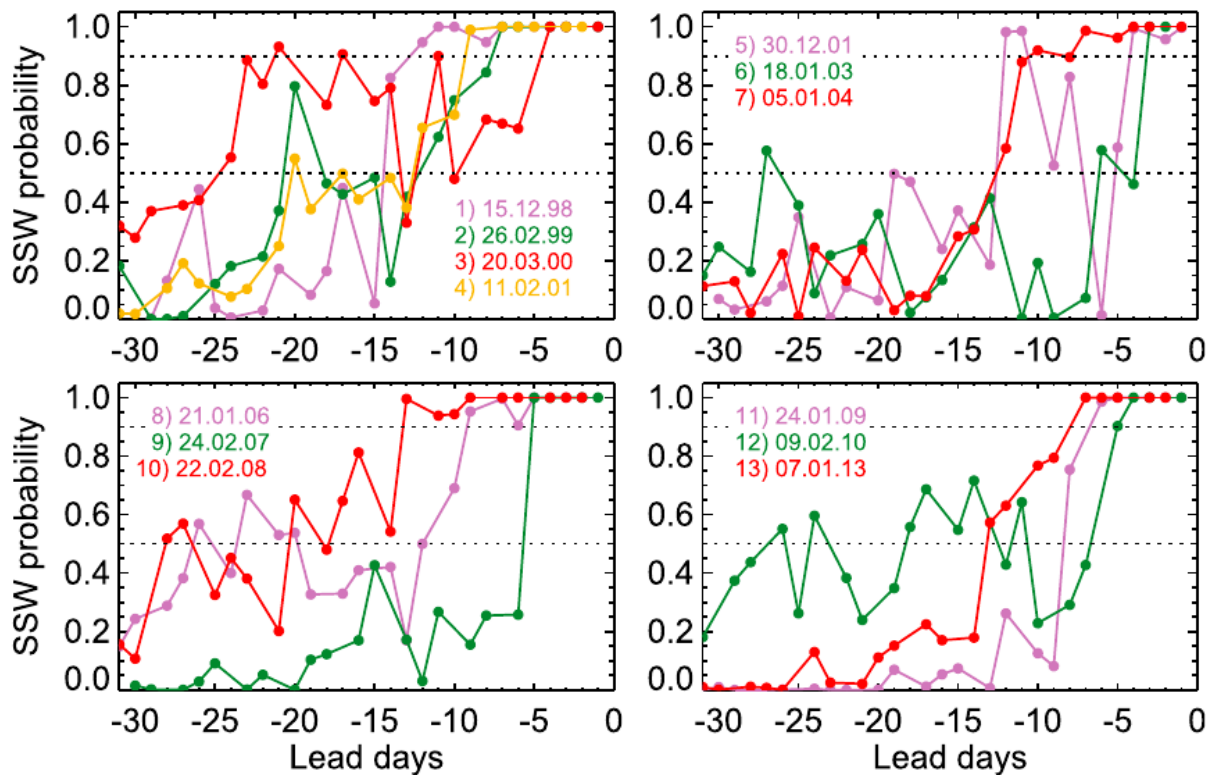


Fig. 2. Forecast probabilities of 13 SSW that occurred on the indicated dates as a function of lead time, based on ensemble hindcasts from the ECMWF monthly forecasting system. Most of the SSWs are predicted between 8 and 12 days lead time with a probability of 0.5–0.9, which is considerably larger than the average frequency of SSW occurrence. (From Karpechko 2018.)

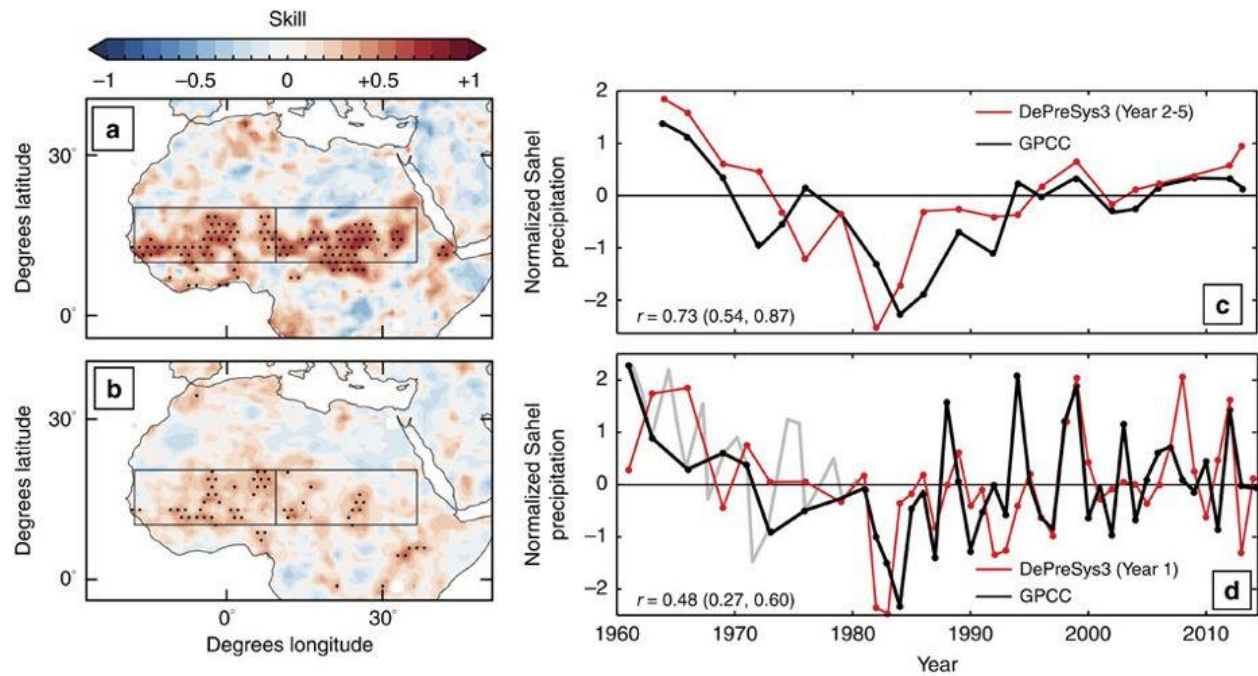


Fig. 3. Skill for predicting linearly detrended Sahel summer rainfall in years 2-5 (upper panels) and year 1 (lower panels) in DePreSys hindcasts. Panels (a)-(b) show spatial distributions of anomaly correlation coefficients with stippling indicating 95% significance. Panels (c)-(d) show time series of normalized predicted and GPCC observed rainfall anomalies in the Sahel region outlined by the boxes in the maps, with correlations and their 5–95% confidence intervals indicated. (From Sheen et al. 2017.)

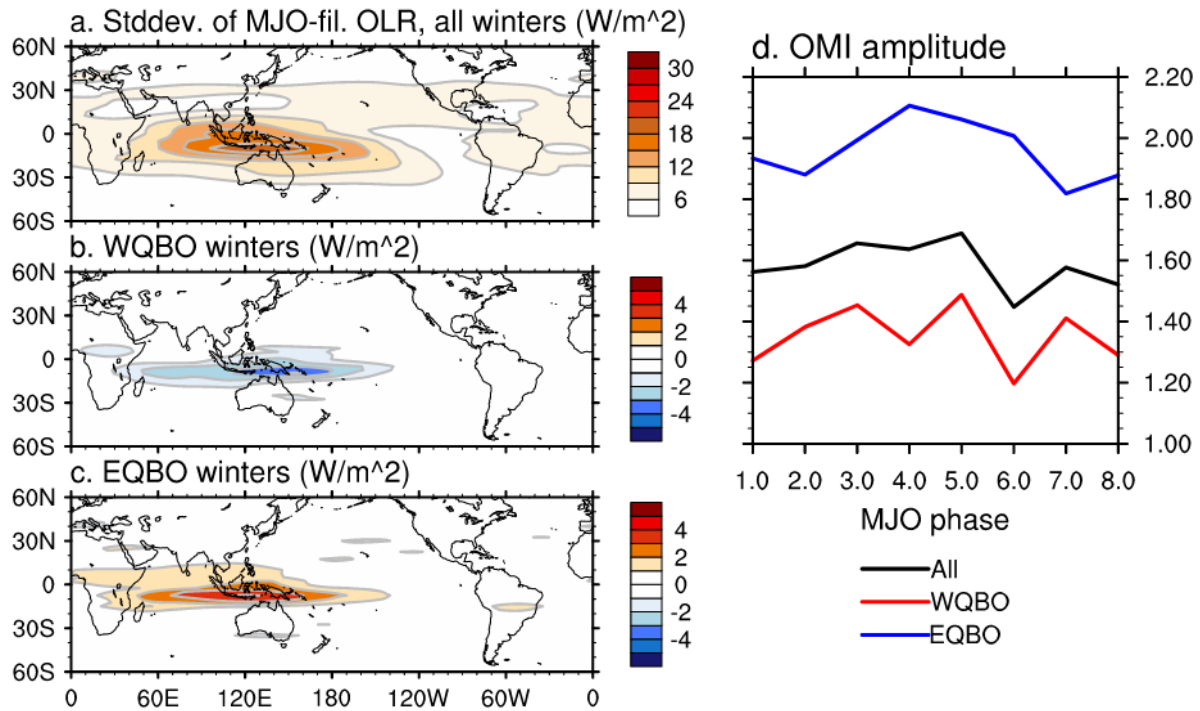


Fig. 4. Influence of QBO phase on MJO amplitude. (a) Standard deviation of wintertime outgoing longwave radiation (OLR), filtered to retain temporal and spatial scales characteristic of the MJO, for all winters in 1979 to 2012. Differences from these values in winters characterized by QBO westerly (WQBO) and easterly (EQBO) phases are shown (b) and (c) respectively. (d) Amplitude of an OLR-based MJO index (OMI) as a function of MJO phase for WQBO, EQBO and all winters. (From Yoo and Son 2016.)

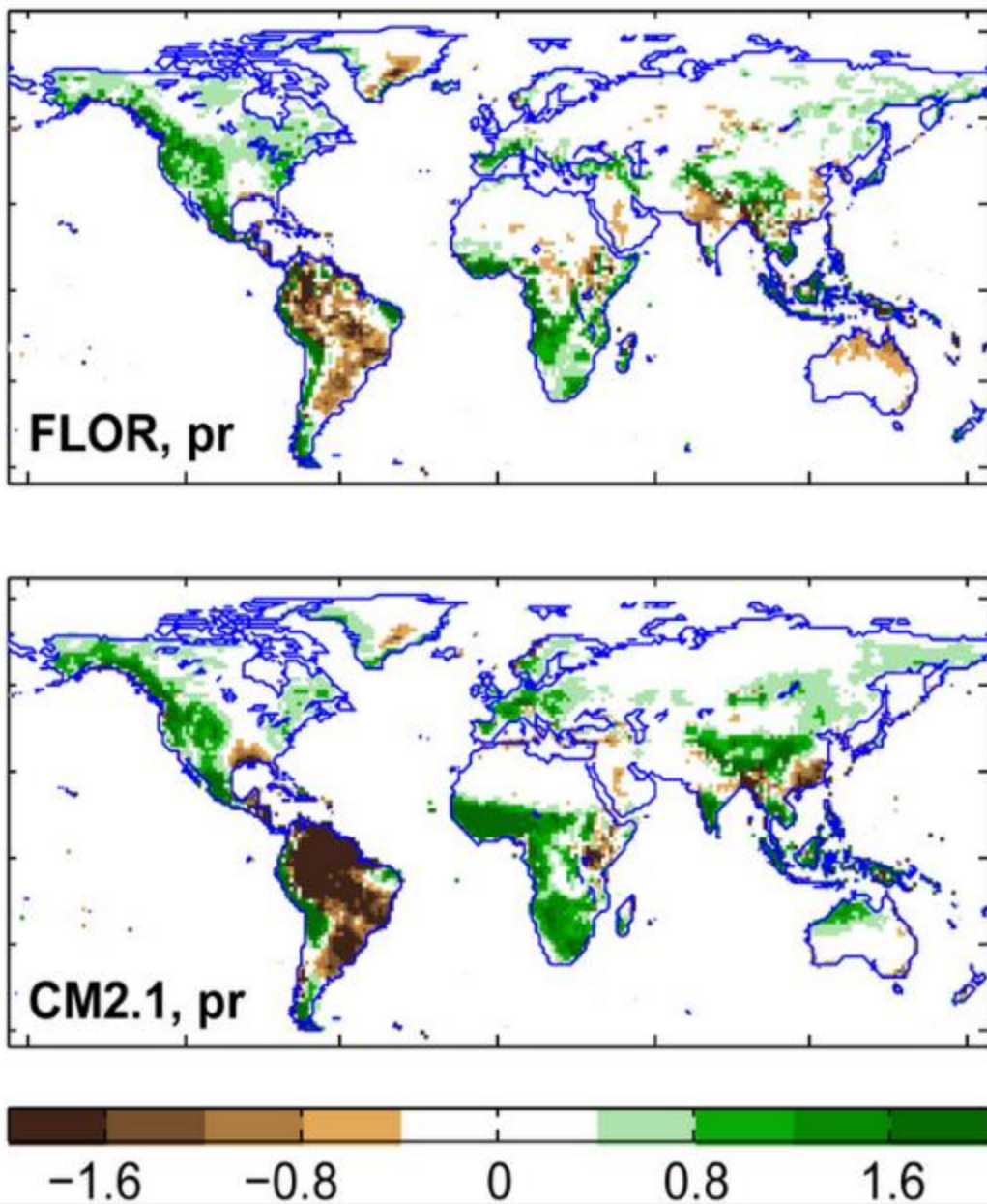


Fig. 5. Impact of resolution on precipitation biases in GFDL seasonal prediction models. Atmospheric resolution is approximately 50 km with 32 levels in FLOR (upper panel), and approximately 200 km with 24 levels in CM2.1 (lower panel), whereas ocean resolution is approximately 100 km in both models. Higher atmospheric resolution in FLOR reduces precipitation biases in numerous regions including much of the tropics. Annual mean biases over land in mm day^{-1} based on 1981-2010 CMAP observations are shown. (After Jia et al. 2015.)

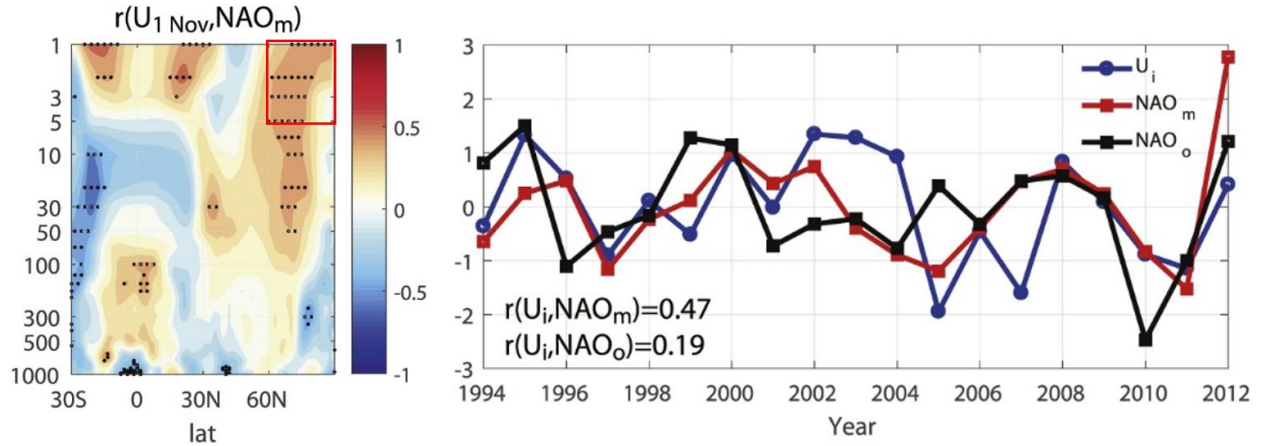


Fig. 6 Connection between stratospheric initial conditions and predicted winter NAO for UK Met Office GloSea5 predictions initialized 1 November 1995-2012. Left: correlation between initial zonal wind anomaly on 1 November and ensemble mean model-predicted surface NAO index (NAO_m) during DJF. Black dots represent values significant at $\alpha = 0.05$ confidence based on one-tailed test, and mean values within the red box define an index U_i . Right: Annual standardized U_i (blue), NAO_m (red) and observed surface NAO index, NAO_o (black). The correlation of U_i with NAO_m , indicated at lower left, is significant at $\alpha = 0.05$ confidence whereas the lower correlation of U_i with NAO_o is not unexpected based on signal to noise considerations and that there is only one realization of observations. The larger correlation of predicted and observed winter NAO values $r(NAO_m, NAO_o)=0.62$ suggests that additional sources of predictability exist. (After Nie et al. 2019.)

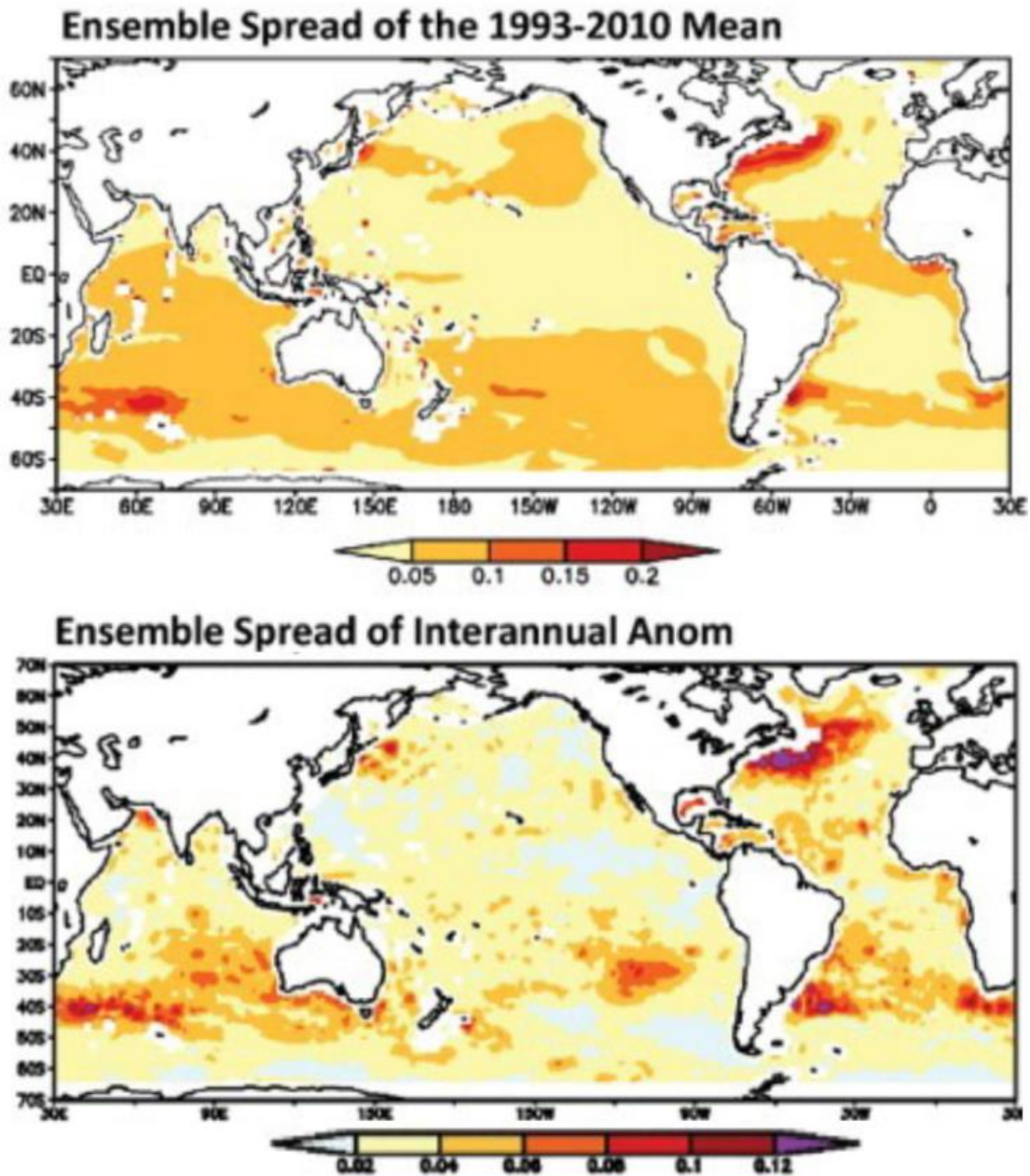


Fig. 7. Consistency across an ensemble of ocean state estimates of depth-averaged salinity over 0–700m, from the Ocean Reanalyses Intercomparison Project. Ensemble standard deviations in both the 1993-2010 means (upper panel) and interannually varying monthly anomalies (lower panel) tend to be largest in eddy active regions such as the Gulf Stream, and less well-observed regions such as the Southern Ocean. These differences across state estimates are reflective of uncertainties in ocean initial conditions. (After Balmaseda et al. 2015.)

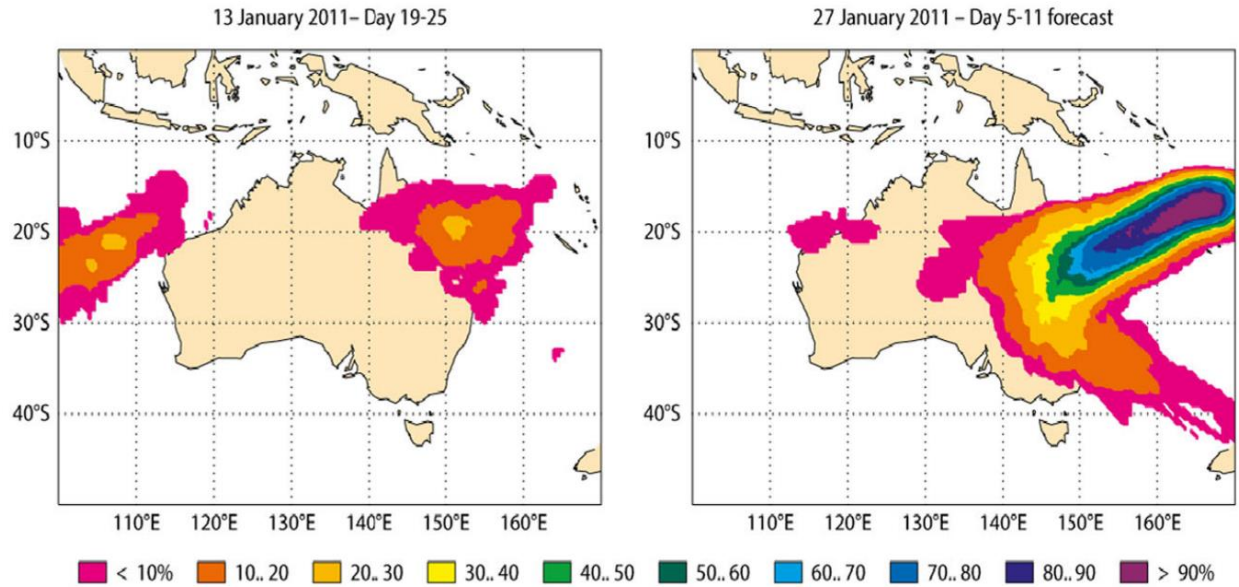


Fig. 8. Elevated probabilities of tropical cyclone occurrence during 31 January to 6 February 2011, based on ECMWF ensemble forecasts forecast starting 13 January with 18 day lead time (left), and 27 January with 4 day lead time (right). Destructive Cyclone Yasi made landfall in northeastern Australia on 3 February 2011 as a destructive category 5 storm. (Adapted from Vitart and Robertson 2018).

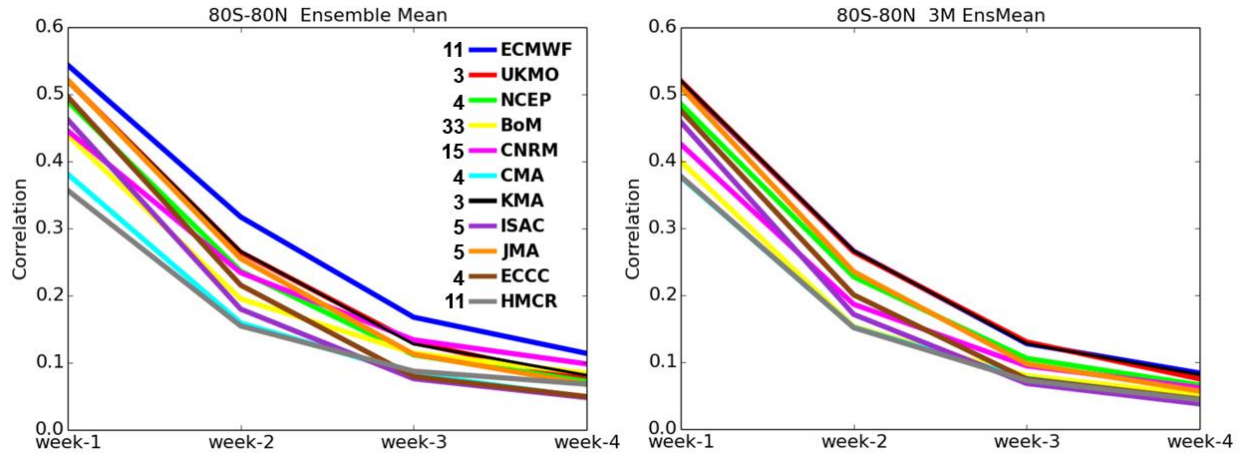


Fig. 9. Global averages of correlations between hindcast and observed precipitation anomalies over the 80°S to 80°N latitudinal band for weeks 1-4 for S2S project models initialized from November to March, 1999–2009. Left: Hindcast quality assessment based on ensemble means using the full ensemble size for each model, as indicated in the figure legend. Right: Hindcast quality assessment based on ensemble means using three ensemble members for each model. The reduced spread of the lines shown in the right panel when ensemble sizes are identical compared to the spread of the lines shown in the left panel demonstrates the value of the use of larger ensembles for subseasonal precipitation forecasting. (Adapted from de Andrade et al. 2019.)

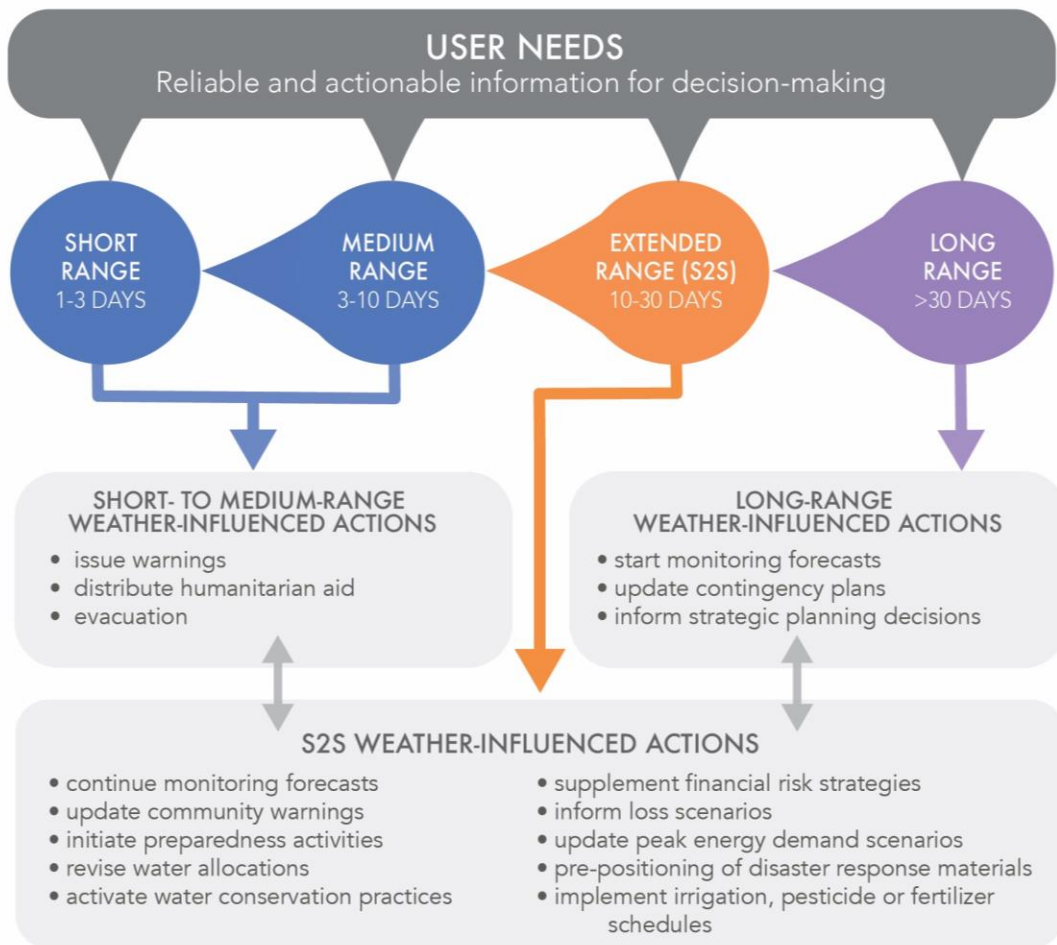


Fig. 10. Schematic illustration of relationships between a S2S forecast range of 10-30 days and other prediction timescales, including examples of actionable information that can enable decision making by various sectors. Indicated actions are examples that are not exclusive to a particular forecast range. (After White et al. 2017.)

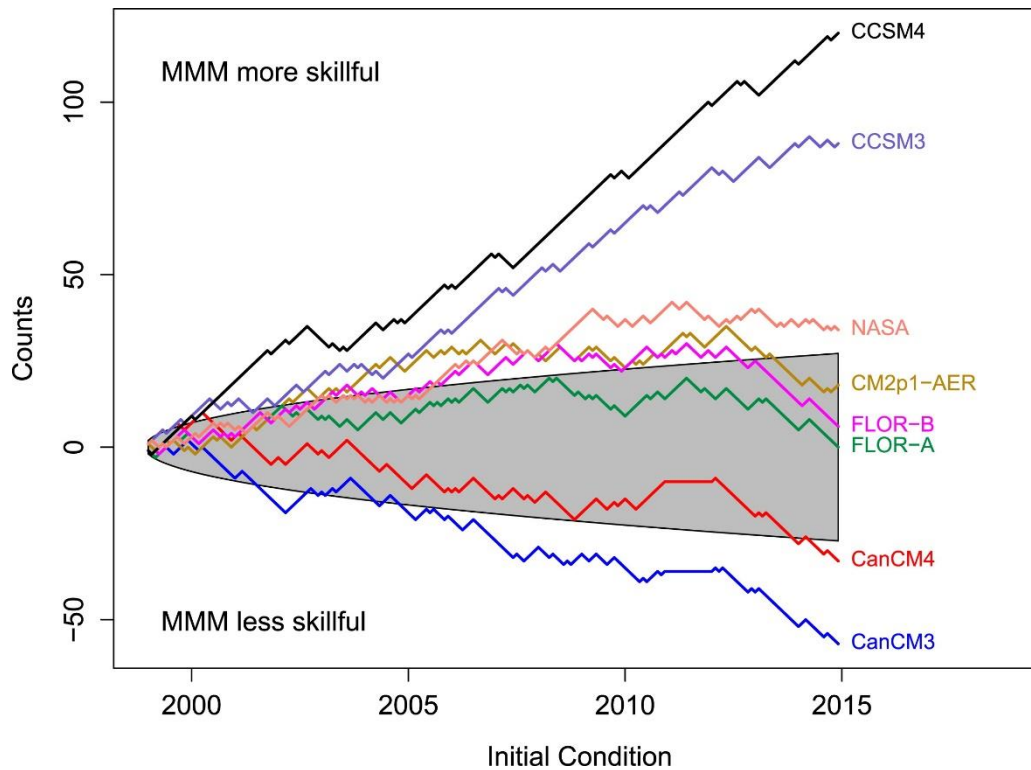


Fig. SB1. Random walk test comparing monthly mean forecasts of the Niño 3.4 index for equatorial Pacific SST at 2.5-month lead, between the multi-model mean (MMM) and individual models in the NMME. Counts (vertical axis) increase by 1 when MMM squared error is smaller than that an individual model (MMM more accurate) and decrease by 1 otherwise (individual model more accurate), and are accumulated forward for all initial months and years (horizontal axis). Accumulated counts above or below the shaded region indicate skill differences according to the squared error metric that are significant with >95% confidence (MMM more skillful above the shaded region and individual model more skillful below). Niño 3.4 anomalies are relative to 1982–98 climatological values, and span each month in 1999–2015. (From DelSole and Tippett 2016.)

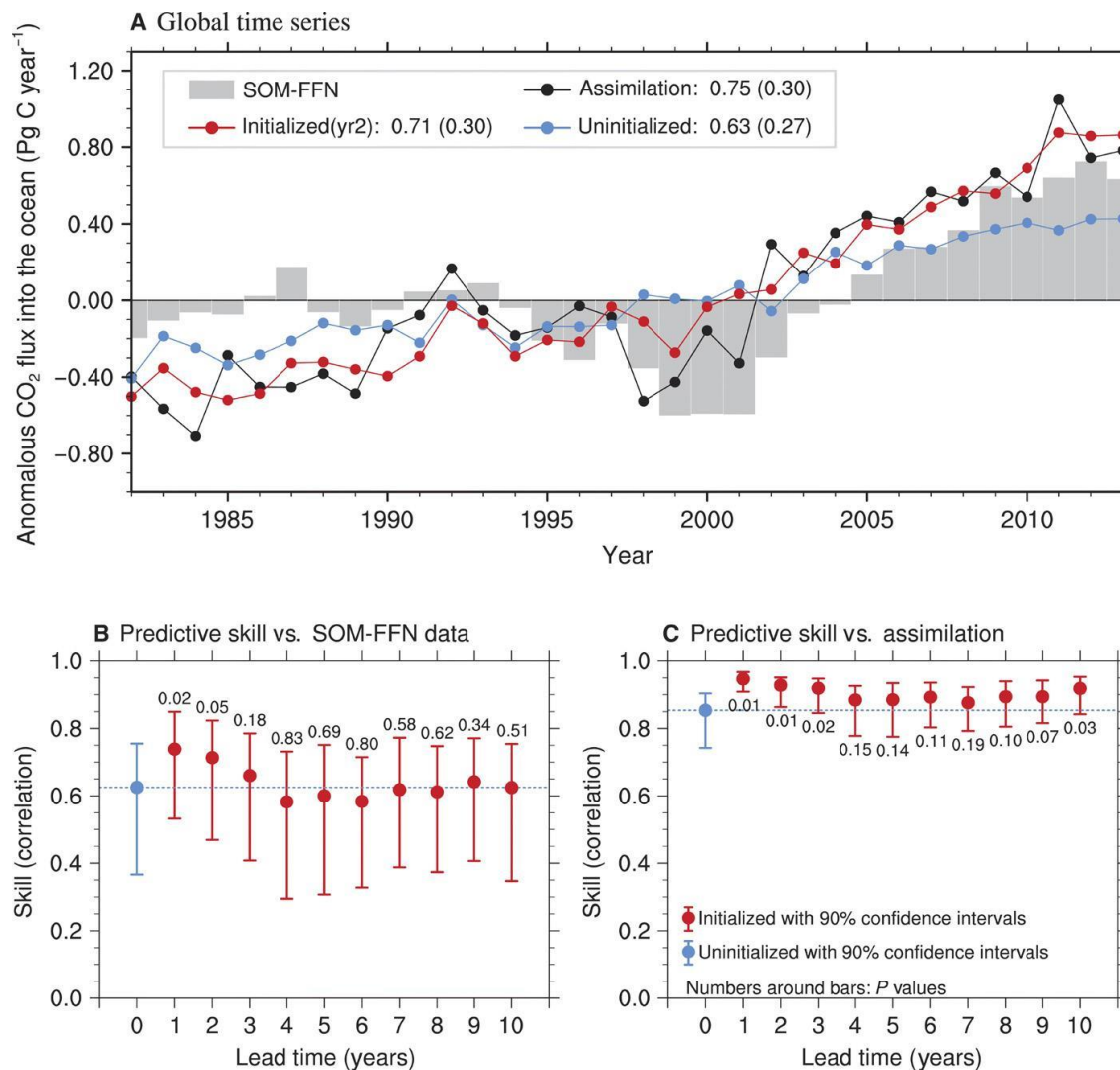


Fig. SB2. Temporal evolution and predictive skill of global CO₂ flux into the ocean from the MPI-ESM-HR decadal prediction system. (A) Annual values of anomalous CO₂ flux into the ocean from data-based estimates (SOM-FFN; gray) and MPI-ESM uninitialized simulations (blue), year 2 of initialized decadal predictions (red) and data-constrained assimilation run (black). Anomaly correlations and root-mean-square errors (in parentheses) verifying against SOM-FFN data are indicated. (B) Anomaly correlation skill for global CO₂ flux into the ocean, verifying against SOM-FFN. The blue dot and dashed line show the uninitialized skill for which lead time is not relevant, and the red dots initialized skill for different forecast years, with 90% confidence intervals and *P* values based on a bootstrap approach indicated. (C) Like (B), but verifying against the assimilation run. (After Li et al. 2019.)